A framework to quantify uncertainty in critical slip distance in rate and state friction model for earthquakes

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Abstract

We arrive at estimates of critical slip distance in the rate and state model for friction evolution using synthetic earthquake data via the Bayesian inference. The conventional solution to the inverse problem is the deterministic parameter values, which may not represent the true value, and quantifying uncertainty in the model parameters increases confidence in the estimation. In this work, the uncertainty in the critical slip distance is estimated by the posterior distribution obtained through the Bayesian inversion.

1 Introduction

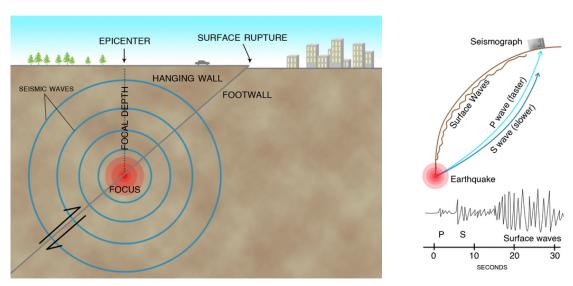


Figure 1: P wave arrives first, followed by the S wave and then by surface waves.

Earthquakes occur as a result of global plate motion. They start out many miles beneath the surface, too deep for us to observe them directly. So they are studied from afar by (1) observing the geological changes at the ground surface, (2) analyzing the symphony of earthquake vibrations recorded on seismographs, and (3) monitoring the tectonic changes in the Earth's crust by surveying it repeatedly, using land survey techniques for many years and now using satellites. An earthquake produces P waves, or compressional waves, that travel faster and reach the seismograph first, and S waves, or shear waves, that are slower (Fig. 1). Both are transmitted within the Earth and are called body waves. Even slower are surface waves that run along the surface of the earth and do a lot of the damage. The earthquake focus is the point

within the Earth where the earthquake originates. The epicenter is a point on the surface directly above the focus. The simplest model for earthquake initiation is to assume that when the stress accumulated in the plates exceeds some failure criterion on a fault plane, an earthquake happens [1]. The groundbreaking work of [2] arrived at the hypothesis that faulting occurs when the resolved shear stress exceeds the internal friction on some plane in the medium leading to fault slip. The quantification of fault slip is achieved using the Rate- and State-dependent Friction (RSF) model for friction evolution, which is considered the gold standard for modeling earthquake cycles (interseismic loading followed by coseismic relaxation) on mature faults [3–7]. It is given by

$$\mu = \mu_0 + A \ln \left(\frac{V}{V_0} \right) + B \ln \left(\frac{V_0 \theta}{d_c} \right),$$

$$\frac{d\theta}{dt} = 1 - \frac{\theta V}{d_c},$$
(1)

where $V = |d\mathbf{d}/dt|$ is the slip rate magnitude, $a = \frac{dV}{dt}$ which we hypothesize is of the same order as recorded by seismograph, μ_0 is the steady-state friction coefficient at the reference slip rate V_0 , A and B are empirical dimensionless constants, θ is the macroscopic variable characterizing state of the surface and d_c is a critical slip distance. The temporal derivatives of friction coefficient are

$$\begin{vmatrix}
\dot{\mu} = \frac{A}{V}\dot{V}, \\
\ddot{\mu} = \frac{A}{V}\ddot{V} - \frac{A}{V^2}(\dot{V})^2
\end{vmatrix}$$
(2)

1.1 Significance of critical slip distance and the lack of calibration thereof

The critical slip distance, d_c , is the distance over which a fault loses or regains its frictional strength after a perturbation in the loading conditions [8]. In principle, it determines the maximum slip acceleration and radiated energy during an earthquake insofar that it influences the magnitude and time scale of the associated stress breakdown process (e.g., fracture energy) [9]. Regardless of the importance, it is paradoxical that the values of d_c reported in the literature range from a few to tens of microns as determined in typical laboratory experiments with bare surfaces and gouge layers [9], to 0.1–5 m as determined in numerical and seismological estimates based on geophysical observations [10], and further to several meters as determined in high-velocity laboratory experiments [11]. Moreover, in most numerical simulations of dynamic rupture propagation with prescribed friction laws, d_c is imposed a priori and its value is often assumed to be constant and uniform on the fault plane. Understanding the physics that controls the critical slip distance and explains the gap between observations from experimental and natural faults is thus one of the crucial problems in both the seismology and laboratory communities [12].

1.2 The agenda for this work

With that in mind, we provide a framework in which synthetic earthquake data is used to quantify uncertainty in critical slip distance. While the resolution and coupled flow and geomechanics [13–23] associated with subsurface activity in the realm of energy technologies and concomitant earthquake quantification is a hot topic, in this work, we focus on the effect of a standard trigonometric perturbation with exponentially decreasing amplitude. In section 2, we explain the spring slider damper idealization to infer the influence of critical slip distance on RSF without recourse to complicated elastodynamic equations. In section 3, we explain the Bayesian inference framework to inversely quantify uncertainty in the estimation of critical slip distance. In section 4, we present conclusions and outlook for future work.

2 The forward model

We first rewrite Eq. (1) as

$$V = V_0 \exp\left(\frac{1}{A} \left(\mu - \mu_0 - B \ln\left(\frac{V_0 \theta}{d_c}\right)\right)\right),$$

$$\dot{\theta} = 1 - \frac{\theta V}{d_c},$$

$$\ddot{\theta} = -\frac{\dot{\theta} V}{d_c}$$

$$(3)$$

As shown in Fig. 2, we model a fault by a slider spring system [24–26]. The slider represents either a fault or a part of the fault that is sliding. The stiffness k represents elastic interactions between the fault patch

and the ductile deeper part of the fault, which is assumed to creep at a constant rate. This simple model assumes that slip, stress, and friction law parameters are uniform on the fault patch.

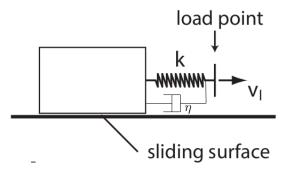


Figure 2: Spring Slider Damper Idealization of Fault Behavior

The friction coefficient of the block is given by

$$\mu = \frac{\tau}{\sigma} = \frac{\tau_l - k\delta - \eta V}{\sigma}$$

where σ is the normal stress, τ the shear stress on the interface, τ_l is the remotely applied stress acting on the fault in the absence of slip, $-k\delta$ is the decrease in stress due to fault slip [1] and η is the radiation damping coefficient [27]. We consider the case of a constant stressing rate $\dot{\tau}_l = kV_l$ where V_l is the load point velocity. The initial stress may be smaller or larger than steady state friction owing to coseismic slip on the fault patch or on adjacent parts of the fault. The expression neglects inertia, and is thus only valid for low slip speed in the interseismic period. The stiffness is a function of the fault length l and elastic modulus E as $k \approx \frac{E}{l}$. With $k' = \frac{E}{l\sigma}$, we get

$$\begin{array}{l}
\dot{\mu} \approx k'(V_l - V) - k''\dot{V}, \\
\ddot{\mu} \approx k'(\dot{V}_l - \dot{V}) - k''\ddot{V}
\end{array} \}$$
(4)

where $k'' = \frac{\eta}{\sigma}$. Once the phenomenological form of $\dot{\mu}$ and $\ddot{\mu}$ is known, we use the following to get \dot{V} and \dot{a} ,

$$\dot{V} = \frac{V}{A} \left(\dot{\mu} - \frac{B}{\theta} \dot{\theta} \right),
\dot{a} = \frac{\dot{V}}{A} \left(\dot{\mu} - \frac{B}{\theta} \dot{\theta} \right) + \frac{V}{A} \left(\ddot{\mu} - \frac{B}{\theta} \ddot{\theta} + \frac{B}{\theta^2} \dot{\theta} \right)$$
(5)

We follow the steps outlined in Algorithm 1 to arrive at the temporal variations of acceleration and fault friction coefficient. We initialize the friction coefficient and state variable and obtain the slip rate and rate of change of the state variable. We then use these values to obtain time derivatives of acceleration and slip rate. These time derivatives are required as we employ the integrated feature of the scientific Python package SciPy [28].

Algorithm 1 Rate and state friction model with radiation damping term

Initialize $\theta = \theta_0$, $\mu = \mu_{ref}$

Use Eq. (3) to get V, $\dot{\theta}$ and $\ddot{\theta}$

Use Eq. (4) to get $\dot{\mu}$ and $\ddot{\mu}$

Use Eq. (5) to get \dot{V} and \dot{a}

3 Bayesian Inversion Framework

The influence of critical slip distance on system response to a load point perturbation of the form

$$V_l = 1 + \exp(-t/10)\sin(10t) \tag{6}$$

is shown in Fig. 3. The code to generate the plots has been given in Appendix A. This code is a part of the GitHub repository https://github.com/karthikncsu/Bayesian-inference-using-earthquake-data.

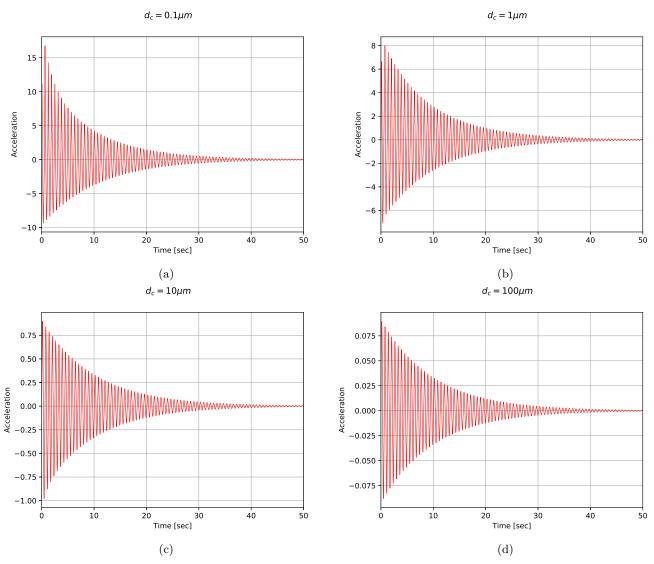


Figure 3: System response for different values of critical slip distance over three orders of magnitude. Units of acceleration are $\mu m/s^2$

The ballpark values are taken from [1] and [27]. Elastic modulus $E = 5 \times 10^{10} \, Pa$, Critical fault length $l = 3 \times 10^{-2} \, m$, Normal stress $\sigma = 200 \times 10^6 \, Pa$, Radiation damping coefficient $\eta = 20 \times 10^6 \, Pa/(m/s)$, A = 0.011 and B = 0.014. The effective stiffness and damping is obtained as

$$k' = \frac{E}{l\sigma} = \frac{5 \times 10^{10}}{3 \times 10^{-2} \times 2 \times 10^{8}} [1/m] \approx 10^{4} [1/m] \equiv 10^{-2} [1/\mu m],$$

$$k'' = \frac{\eta}{\sigma} = \frac{2 \times 10^{7}}{2 \times 10^{8}} = 0.1 [s/m] \approx 1 \times 10^{-7} [s/\mu m]$$

In an inverse problem, the acceleration response of the model is known and the goal is to find the parameter, critical slip distance parameter (d_c) . To define the inverse problem, considered the relationship between acceleration $(a_i(t))$ and the model response by the following statistical model

$$a(t_i) = f(t_i, \theta, \mu, A, B, d_c) + \epsilon_i \tag{7}$$

where ϵ_i is the error in the statistical model. Here the $a(t_i)$ and ϵ_i are the random variables. The earthquake data over time $a(t_1), ..., a(t_n)$ are the n observations for $a(t_i)$ and $f(A, t_i, \theta, \mu, A, B, d_c)$ is the acceleration response of the model over time obtained using the Algorithm 1. The goal of the inverse problem is to determine the model parameter (d_c) from the Eq.(7) and conventional method to determine the model parameter that mimizes the norm of the errors using the least squares fit solution as shown below

$$d_{c,0} = \arg\min_{d_c} \sum_{i=1}^{n} (\epsilon_i)^2 = \arg\min_{d_c} \sum_{i=1}^{n} (a(t_i) - f(t_i, \theta, \mu, A, B, d_c))^2$$
(8)

The critical slip distance parameter (d_c) obtained using the least-squares fit solution, Eq.(8) is deterministic value. The values estimated using the least square fit are not the true values due to inherent noise in the data and in most cases, the noise in the data makes it difficult to find the true value. Instead, finding a probability distribution for the model parameters encompasses the true model parameter values and increases the confidence in the prediction. Using the Bayes theorem [29], the distribution for the model parameters is given by the posterior distribution

$$\pi(d_c|a(t_1),...,a(t_n)) = \frac{\pi(a(t_1),...,a(t_n)|d_c)\pi_0(d_c)}{\int_{d_c} \pi(a(t_1),...,a(t_n)|d_c)\pi_0(d_c)dd_c}$$
(9)

Here $\pi(d_c|a(t_1),...,a(t_n))$ is the posterior, $\pi(a(t_1),...,a(t_n)|d_c)$ is the likelihood and $\pi_0(d_c)$ is the prior distribution for the model parameters. Assuming the $\epsilon_i \sim N(0,\sigma^2)$ as unbiased, independent and identical normal distribution with standard deviation σ , the likelihood function is expressed as

$$\pi(a(t_1), ..., a(t_n)|d_c) = \prod_{i=1}^n \pi(a(t_i)|d_c) = \prod_{i=1}^n \frac{1}{\sigma\sqrt{2\pi}} e^{-\frac{1}{2}\left(\frac{a(t_i) - f(V, t_i, \theta, \mu, A, B, d_c)}{\sigma}\right)^2}$$
(10)

In above equation, the $f(V, t_i, \theta, \mu, A, B)$ is calculated using the forward problem given by algorithm 1. The information of the model parameters can be included in the posterior distribution through the prior, $\pi_0(d_c)$. In this study, the prior is assumed to be uniform distribution and the prior is a constant value inside the uniform distribution limits.

Algorithm 2 Bayesian inference

Input data: earthquake data, $a(t_1), ..., a(t_n)$

Generate grid d_c

Use Algorithm 1 to get RHS of Eq. (7)

Use (10) to get $\pi(a(t_1),...,a(t_n)|d_c)$ for each grid point.

Integrate Eq. (9) to obtain the posterior distribution

The goal of the inverse problem is to calculate the posterior distribution Eq.(9), which represents the uncertainty in the critical slip distance parameter (d_c) due to the noise in the earthquake data. Direct evaluation of the posterior distribution using quadrature rules is expensive and often requires adaptive methods to find the posterior distribution. Alternatively, sampling methods like Markov chain Monte Carlo (MCMC) methods [29–31] can be used to generate samples from the posterior distribution.

4 Conclusions and Outlook

This work presents a framework to inversely quantify uncertainty in the critical slip distance parameter of the rate and state friction (RSF) model via the Bayesian inference using the earthquake data. The forward model is to determine the acceleration, using the RSF model, for the given model parameters. In case of an inverse problem, the acceleration data is known and the goal is to find the model parameters. Using conventional methods such as least-squares methods, a deterministic value of the critical slip distance parameter can be obtained from the inverse problem. However, the deterministic parameter value estimated

using the conventional methods does not represent the true values due to the noise in the earthquake data, and quantifying uncertainty in the model parameters increases the confidence in the prediction. The uncertainty in the model parameter is estimated by the posterior distribution obtained from the Bayes theorem. The future work will be to demonstrate a simulation to quantify uncertainty in the critical slip distance parameter using the earthquake data via sampling methods such as Markov chain Monte Carlo.

A Python Program to Generate Forward Model Response

```
1 from __future__ import division
2 from scipy.misc import derivative
3 import numpy as np
4 from scipy import integrate
5 import matplotlib.pyplot as plt
6 from math import exp, log, pi, sin, cos
8 fig = plt.figure()
9 fig.suptitle('$d_c=100 \mu m$')
global Dc, V_ref, delta_t, t_start, t_final, amp, num_steps, k
12 Dc = 100
13 V_ref = 1
t_start = 0.0
t_final = 50.
16 \text{ delta\_t} = 1e-2
num_steps = int((t_final-t_start)/delta_t)
18
  def friction(t,y):
19
20
      a = 0.011
      b = 0.014
21
      kprime = 1e-2*10/Dc # inversely prop to Dc
22
       mu ref = 0.6
23
      k1 = 1e-7 \# radiation damping term
24
25
       V_1 = V_{ref}*(1+exp(-t/10)*sin(10*t))
26
2.7
       # Just to help readability
28
29
       #y[0] is mu (friction)
       #y[1] is theta
30
       #y[2] is velocity
31
32
      n = len(y)
33
       dydt = np.zeros((n,1))
34
35
       # compute v
36
       temp_ = V_ref * y[1] / Dc
37
       temp = 1/a*(y[0] - mu_ref - b * log(temp_))
38
       v = V_ref * exp(temp)
39
40
       # time derivative of theta
41
       dydt[1] = 1. - v * y[1] / Dc
42
43
       # double derivative of theta
44
       ddtheta = - dydt[1]*v/ Dc
45
46
       # time derivative of mu
47
       dydt[0] = kprime*V_l - kprime*v
48
49
       # time derivative of velocity
50
       dydt[2] = v/a*(dydt[0] - b/y[1]*dydt[1])
       # radiation damping
53
       dydt[0] = dydt[0] - k1*dydt[2]
54
55
       dydt[2] = v/a*(dydt[0] - b/y[1]*dydt[1])
56
```

```
return dydt
57
58
59 r = integrate.ode(friction).set_integrator('vode', order=5,max_step=0.001,method='bdf',
      atol=1e-10, rtol=1e-6)
60
61 # Initial conditions
mu_t_zero = 0.6
mu_ref = 0.6
64 theta_t_zero = Dc/V_ref
65 v = V_ref
66 r.set_initial_value([mu_t_zero, theta_t_zero, V_ref], t_start)
68 # Create arrays to store trajectory
69 t = np.zeros((num_steps,1))
70 mu = np.zeros((num_steps,1))
71 theta = np.zeros((num_steps,1))
72 velocity = np.zeros((num_steps,1))
73 acc = np.zeros((num_steps,1))
74 t[0] = t_start
75 mu[0] = mu_ref
76 theta[0] = theta_t_zero
77 \text{ velocity}[0] = v
78 \ acc[0] = 0
79
80 # Integrate the ODE(s) across each delta_t timestep
81 k = 1
82 while r.successful() and k < num_steps:</pre>
       #integrate.ode.set_f_params(r,velocity,k)
83
84
       r.integrate(r.t + delta_t)
85
       # Store the results to plot later
86
87
       t[k] = r.t
       mu[k] = r.y[0]
88
       theta[k] = r.y[1]
89
       velocity[k] = r.y[2]
90
       acc[k] = (velocity[k]-velocity[k-1])/delta_t
91
92
93
94 # Make some plots
95
96 \text{ ax4} = \text{plt.subplot}(111)
97 ax4.plot(t, acc, 'r', linewidth=0.5)
98 ax4.set_xlim(t_start, t_final)
99 ax4.set_xlabel('Time [sec]')
ax4.set_ylabel('Acceleration')
101
   ax4.grid('on')
103 plt.show()
```

Listing 1: Python example

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