



# Fast Probabilistic Seismic Hazard Analysis through Adaptive Importance Sampling

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## ABSTRACT

Probabilistic Seismic Hazard Analysis (PSHA) relies on two widely utilized approaches with high computational demands: (a) Riemann sum and (b) conventional Monte Carlo (MC) integration. The first requires sufficiently fine slices across magnitude, distance, and ground motion, and the second requires extensive synthetic earthquake catalogs to compute seismic hazards accurately. These approaches are notably resource-intensive for low-probability seismic hazards, e.g., up to  $10^8$  MC samples for a hazard with  $10^{-4}$  probability to achieve coefficient of variation (COV) of 1%. Here, we present a novel framework to compute hazard and deaggregation with unprecedented computational efficiency. We formulate Adaptive Importance Sampling (AIS) PSHA to approximate optimal important sampling (IS) distributions and dramatically reduce the size of synthetic earthquake catalogs (i.e., number of MC samples) to estimate hazards. We evaluate the effectiveness and reliability of our proposed method using comprehensive test problems from the Pacific Earthquake Engineering Research Center (PEER) for PSHA benchmarks, encompassing various seismic source types, including areal, fault, and combined ones. Our findings indicate that this novel approach significantly outpaces Riemann sum and traditional MC methods with computations up to  $>10^5$  and  $7.8 \times 10^3$  times faster, respectively, while maintaining a standard deviation of the estimate below 2%. Moreover, we show theoretically that optimal IS distributions are equivalent to hazard deaggregation distributions. Empirically, we show our approximated optimal IS and the deaggregation distributions are closely alike, e.g., with a Kolmogorov–Smirnov statistic between 0.017 and 0.113. We developed our methodology to have broad applicability in PSHA practices, especially in cases requiring extensive computational resources to navigate numerous logic tree scenarios addressing epistemic uncertainty.

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## KEY POINTS

- We introduce an adaptive importance sampling algorithm for Probabilistic Seismic Hazard Analysis (PSHA).
- The proposed algorithm is up to  $10^3$  and  $10^5$  times more efficient than conventional Monte Carlo and Riemann sum.
- The algorithm also facilitates the straightforward implementation of hazard deaggregation.

### Supplemental Material

## INTRODUCTION

Probabilistic Seismic Hazard Analysis (PSHA) has become a foundational method for determining seismic design levels and conducting regional seismic risk analyses since its first inception (Cornell, 1968; U. S. Nuclear Regulatory Commission, 2007; McGuire, 2008; ASCE, 2022). The development of PSHA was driven by the necessity for a probabilistic framework to quantify seismic hazards, acknowledging the unpredictable nature of earthquake location, magnitudes, and ground motion intensities (Esteva, 1967; Cornell, 1968). In PSHA, earthquake location, magnitude, and ground motions are treated as random variables, facilitating the computation of annual exceedance probabilities at various ground motion intensities. Integrating hazard with fragility curves further enables to determine the annual probabilities of structural failures, thereby underscoring the methodology's critical role in risk assessment (Kennedy et al., 1980; Ceferino et al., 2020; Silva et al., 2020; Baker et al., 2021; Papadopoulos and Bazzurro, 2021; Arora and Ceferino, 2023).

Since we cannot solve PSHA analytically due to the complexity in seismic source and ground motion models, numerous researchers have developed computer softwares for PSHA computation (Cornell, 1968; McGuire, 1976; Field et al., 2003; Ordaz et al., 2013; Pagani et al., 2014). According to the PSHA computer code verification project (Thomas et al., 2010; Hale et al., 2018), existing software primarily employs Riemann summation for numerical integration of PSHA, which partitions the earthquake magnitude, location, and ground motion random variable into fine grids to approximate the actual integration. The Riemann summation offers robust PSHA integration with sufficiently dense grids. However, this method generally incurs a significant computational load exponentially increasing with the number of grids and dimensions in multi-dimensional integrations (Philippe and Robert, 2001). Furthermore, the results are highly sensitive to the chosen grid design (e.g., the initial point of the grid, grid spacing), leading to significant deviations from one software to another, especially for low exceedance probabilities (Thomas et al., 2010; Hale et al., 2018).

Alternatively, other softwares adopted Monte-Carlo (MC) integration for PSHA (Assatourians and Atkinson, 2013, 2019). MC integration calculates exceedance probabilities by generating synthetic earthquake catalogs based on the seismic source and ground motion models to evaluate the recurrence of various ground motion intensities (Musson, 2000). MC integration's primary advantage lies in its straightforward concept compared to Riemann summation, without the need to divide the integration range into small slices (Musson, 2000; Dick et al., 2013). Nevertheless, MC framework requires a substantially-

46 long synthetic catalog to accurately estimate hazard from rare events, making the computation costly, especially for large  
47 ground motions with low exceedance probabilities, e.g.,  $p < 10^{-4}$  /yr (Kroese et al., 2014).

48 Importance Sampling (IS) can offer a solution to this rare event simulation (Tokdar and Kass, 2010). IS was initially intro-  
49 duced in statistical physics (Hammersley and Morton, 1954) to improve the computational efficiency of rare event simulation  
50 that would otherwise require a large sample size with conventional MC. IS relies on identifying an appropriate probability  
51 distribution (“IS distribution”) to explore low-probability spaces effectively. Researchers use the IS distribution to sample  
52 rare events with a higher likelihood than conventional MC and then correct their frequency through weights, significantly  
53 reducing the number of samples to compute low probabilities (Robert et al., 1999). However, finding such an appropriate dis-  
54 tribution can be challenging because there is no optimal sampling density that is universally applicable; rather, the proper  
55 selection of sampling density depends on the problem being solved. Thus, many numerical experiments are often conducted  
56 first through trial-and-error to identify IS distributions, which is still computationally expensive.

57 In regional seismic risk analysis, numerous studies have been conducted to sample hazard-consistent earthquake ground  
58 motions (Crowley and Bommer, 2006; Kiremidjian et al., 2007; Jayaram and Baker, 2010; Han and Davidson, 2012; Manzour  
59 et al., 2016; Christou et al., 2018; Kavvada et al., 2022). Kiremidjian et al. (2007) first introduced IS distributions that samples  
60 large-magnitude earthquakes with a high probability to reduce the computational burden of seismic hazard and risk analyses.  
61 Jayaram and Baker (2010) expanded the approach by defining IS distributions to sample high-intensity ground motions.  
62 However, Jayaram and Baker (2010) highlighted the computational challenges to identify an effective IS distribution and  
63 ended up using K-mean clustering to reduce the number of ground motion samples. Rahimi and Mahsuli (2019) applied  
64 system reliability methods to calculate PSHA. They selected the IS sampling density as the normal distribution centered at  
65 the “design point” derived from the first- and second-order reliability method.

66 To find effective IS distributions, computational statisticians have developed a general framework to find them through  
67 iterative algorithms denominated “Adaptive Importance Sampling (AIS)” (Bugallo et al., 2017). AIS algorithms leverage the  
68 fact that the optimal density is proportional to the integrand of the problem to find the sampling density which minimizes  
69 the variance of the estimate. Various AIS algorithms have been introduced, e.g., cross-entropy based AIS, Vegas, Divonne,  
70 and Miser (Lepage, 1978; Friedman and Wright, 1981; Rubinstein, 1997; Press and Farrar, 1990; Rubinstein and Kroese, 2004;  
71 Bugallo et al., 2017). The application of AIS is widely adopted to solve the integration with high dimensions, such as in the  
72 field of statistical physics, finance, reliability engineering, and signal processing (Au and Beck, 2001; Kappen and Ruiz, 2016;  
73 Nieto and Ruiz, 2016; Bugallo et al., 2017). Although previous studies suggest the use of importance sampling (IS) for PSHA  
74 calculation (Jayaram and Baker, 2010; Rahimi and Mahsuli, 2019), no research has been published regarding the use of AIS  
75 for this purpose. AIS can provide a general methodology for identifying an appropriate IS distribution for seismic hazards,  
76 eliminating computationally expensive experiments in the regular IS approach, whose efficiency typically depends on the  
77 researcher’s experience in the field.

78 Many AIS algorithms have been published (Lepage, 1978; Friedman and Wright, 1981; Rubinstein, 1997; Press and Farrar,  
79 1990; Rubinstein and Kroese, 2004; Bugallo et al., 2017). However, we study the VEGAS algorithm for PSHA to leverage its  
80 straightforward mathematical framework and fast convergence. Thus, this study introduces a novel computational method  
81 for PSHA curve calculation using Adaptive Importance Sampling (AIS) VEGAS algorithm (Lepage, 1978, 2021). This paper  
82 also shows that AIS facilitates hazard deaggregation, the relative contribution of each random variables to the overall hazard  
83 (Bazzurro and Cornell, 1999), because the optimal IS density resembles the contributions of each variable—magnitude,  
84 distance, and ground motion—to the hazard. Thus, our proposed framework also enhances the computational efficiency of  
85 deaggregation estimates, which otherwise would require additional memory resources and complexity in the computer code  
86 with traditional calculation methods. We explore three key aspects of the method: 1) the enhancement of computational  
87 efficiency that the AIS algorithm offers over traditional methods; 2) the accuracy of the estimates provided by this approach;  
88 and 3) the process of obtaining hazard deaggregation through AIS. We present the theoretical background of AIS PSHA and  
89 validate the method through numerical examples.

## 90 MATHEMATICAL FORMULATION

### 91 Probabilistic Seismic Hazard Analysis (PSHA)

92 At a site of interest, the annual frequency of ground motion exceedance from a single source can be calculated as:

$$\lambda(X > a) = \nu \int_{\varepsilon_{\min}}^{\varepsilon_{\max}} \int_{r_{\min}}^{r_{\max}} \int_{m_{\min}}^{m_{\max}} I(X > a | m, r, \varepsilon) f_{M,R,\varepsilon}(m, r, \varepsilon) dm dr d\varepsilon \quad (1)$$

93 , where  $\lambda(X > a)$  is annual rate that ground motion,  $X$ , exceeds the target ground motion intensity,  $a$ , e.g., peak ground  
94 acceleration.  $\nu$  is the annual rate of earthquake occurrence greater than  $m_{\min}$  from the source,  $m$  is the earthquake magnitude,  
95  $m_{\min}$  and  $m_{\max}$  are minimum and maximum magnitudes considered for the source,  $r$  is the source-to-site distance,  $r_{\min}$  and  
96  $r_{\max}$  are minimum and maximum source-site distances,  $\varepsilon$  is a standard normal random variable for generating earthquake  
97 ground motion,  $\varepsilon_{\min}$  and  $\varepsilon_{\max}$  are minimum and maximum  $\varepsilon$  (generally,  $\varepsilon_{\max} \geq 6$  and  $\varepsilon_{\min} \leq -6$ ; Bommer and Abrahamson  
98 (2006)),  $f_{M,R,\varepsilon}(m, r, \varepsilon)$  is joint probability density function (PDF) of  $m$ ,  $r$ , and  $\varepsilon$ ,  $I(X > a | m, r, \varepsilon)$  is indicator function that  
99 takes 1 when  $X > a$ , otherwise, 0. The ground motion  $X$  given  $M$ ,  $R$ , and  $\mathcal{E}$  is generally calculated using ground motion  
100 models (Bozorgnia et al., 2014; Goulet et al., 2021). The models usually assume the log-normal distribution for the ground  
101 motion given explanatory variables such as  $M$  and  $R$ . Naturally, these models provide the mean and standard deviation of  
102 logarithmic ground motion. Thus, the random ground motion is calculated as:

$$\log X = \mu(M, R) + \mathcal{E}\sigma(M, R) \quad (2)$$

103 , where  $\mu$  and  $\sigma$  are mean and standard deviation of logarithmic earthquake ground motion. By taking exponential on both  
104 sides of Eq. (2), the ground motion  $X$  can be calculated as

$$X = e^{\mu(M,R) + \varepsilon\sigma(M,R)}$$

105 If we assume that the ground motion random variable  $\varepsilon$  is independent with respect to the  $m$  and  $r$  (McGuire, 1995), Eq. (1)  
 106 can be modified as

$$\begin{aligned} \lambda(X > a) &= \nu \int_{\varepsilon_{\min}}^{\varepsilon_{\max}} \int_{r_{\min}}^{r_{\max}} \int_{m_{\min}}^{m_{\max}} I(x > a | m, r, \varepsilon) f_{\varepsilon}(\varepsilon) f_{M,R}(m, r) dm dr d\varepsilon \\ &= \nu \int_{\varepsilon_{\min}}^{\varepsilon_{\max}} \int_{R_{\min}}^{R_{\max}} \int_{M_{\min}}^{M_{\max}} I(x > a | m, r, \varepsilon) f_{\varepsilon}(\varepsilon) f_{R|M}(r | m) f_M(m) dm dr d\varepsilon \end{aligned} \quad (3)$$

107 , where  $f_M(m)$ ,  $f_{R|M}(r | m)$ ,  $f_{\varepsilon}(\varepsilon)$  are PDF of  $m$ ,  $r$  given  $m$ , and  $\varepsilon$ . Under point source assumption, the distance  $r$  and magnitude  
 108  $m$  become independent random variables. Thus, the seismic hazard is given by

$$\lambda(X > a) = \nu \int_{\varepsilon_{\min}}^{\varepsilon_{\max}} \int_{r_{\min}}^{r_{\max}} \int_{m_{\min}}^{m_{\max}} I(x > a | m, r, \varepsilon) f_{\varepsilon}(\varepsilon) f_R(r) f_M(m) dm dr d\varepsilon \quad (4)$$

109 The total seismic hazard from multiple seismic sources (e.g., different faults) is the sum of each. Thus,

$$\Lambda(X > a) = \sum_{i=1}^n \lambda_i(X > a) \quad (5)$$

110 , where  $\Lambda(X > a)$  is the total annual frequency of exceedance of ground motion,  $a$ ,  $i$  is index for seismic sources, and  $n$  is the  
 111 total number of seismic sources. Under the assumption of Poisson process, the annual probability of exceedance,  $\phi(X > a)$ ,  
 112 can be converted from annual frequency of exceedance ( $\Lambda$ ; Eq. (5)) as

$$\phi(X > a) = 1 - e^{-\Lambda(X > a)} \quad (6)$$

### 113 Hazard deaggregation

114 We can also deaggregate the total hazard (Eq. (6)) to better understand the earthquakes that contribute most to the hazard.  
 115 Deaggregation is also used to develop the select seismic records (e.g., from the earthquakes that contribute most to the hazard)  
 116 and conduct non-linear time-history analyses for the design of many critical buildings (Bazzurro and Cornell, 1999; U. S.  
 117 Nuclear Regulatory Commission, 2007). Mathematically, deaggregation of the hazard is the joint probability distribution of  
 118 the  $m$ ,  $r$ , and  $\varepsilon$  conditional on different levels of hazards  $a$  to quantify the contributions of each component. The deaggregation  
 119 of PSHA can be formulated using Bayes' theorem as

$$\begin{aligned}
P(m, r, \varepsilon|X > a) &= \frac{P(X > a \cap m, r, \varepsilon)}{P(X > a)} \\
&= \frac{P(X > a|m, r, \varepsilon)P(m, r, \varepsilon)}{P(X > a)} \\
&= \frac{P(X > a|m, r, \varepsilon)P(m, r, \varepsilon)}{\sum P(X > a|m, r, \varepsilon)P(m, r, \varepsilon)} \\
&= \frac{I(X > a|m, r, \varepsilon)P(m, r, \varepsilon)}{\sum I(X > a|m, r, \varepsilon)P(m, r, \varepsilon)} \tag{7}
\end{aligned}$$

120 , where  $P(X > a|m, r, \varepsilon)$  is the probability of ground motion  $X$  is greater than  $a$  given  $m, r$ , and  $\varepsilon$ ,  $P(m, r, \varepsilon)$  is joint probability  
121 of  $m, r$ , and  $\varepsilon$ , and  $P(X > a)$  is the total probability that the ground motion is greater than  $a$ , which is the summation of  
122  $P(X > a|m, r, \varepsilon)P(m, r, \varepsilon)$  over all  $m, r$ , and  $\varepsilon$ . Note that  $P(X > a|m, r, \varepsilon)$  can be expressed as an indicator function,  $I(X >$   
123  $a|m, r, \varepsilon)$ , because the probability of ground motion  $X$  greater than  $a$  can only be 1 or 0 given  $m, r$ , and  $\varepsilon$  (see Eq. (2)).

124 By replacing probability mass function,  $P(m, r, \varepsilon)$ , with probability density function,  $f_{M,R,\varepsilon}(m, r, \varepsilon)$ , and change the  
125 summation into integration, Eq. (7) can be expressed as:

$$f(m, r, \varepsilon|X > a) = \frac{I(X > a|m, r, \varepsilon)f_{M,R,\varepsilon}(m, r, \varepsilon)}{\iiint I(X > a|m, r, \varepsilon)f_{M,R,\varepsilon}(m, r, \varepsilon)dmdr d\varepsilon} \tag{8}$$

126 By Eq. (1), the denominator of Eq. (8) equals  $\lambda/\nu$ . Thus,

$$f(m, r, \varepsilon|X > a) = \nu \cdot \frac{I(X > a|m, r, \varepsilon)f_{M,R,\varepsilon}(m, r, \varepsilon)}{\lambda} \tag{9}$$

127 This equation shows that the contribution of specific  $m, r$ , and  $\varepsilon$  can be represented by the ratio of the partial sum of the  
128 given  $m, r$ , and  $\varepsilon$  to the total hazard. Note that  $\lambda$  and  $\nu$  are constant. Thus, the hazard deaggregation,  $f(m, r, \varepsilon|x > a)$ , is  
129 proportional to  $I(x > a|m, r, \varepsilon)f(m, r, \varepsilon)$ :

$$f(m, r, \varepsilon|x > a) \propto I(x > a|m, r, \varepsilon)f(m, r, \varepsilon)$$

## 130 CURRENT NUMERICAL SOLUTIONS

131 **The Riemann summation** This method computes PSHA curves by summing the areas of partitioned  $(m, r, \varepsilon)$  cuboids. The  
132 Riemann summation for Eq. (1) can be expressed as:

$$\lambda(X > a) = \nu \sum_{k=1}^{N_\varepsilon} \sum_{j=1}^{N_r} \sum_{i=1}^{N_m} I(x > a|m_i, r_j, \varepsilon_k) f_{M,R,\varepsilon}(m_i, r_j, \varepsilon_k) \Delta m \Delta r \Delta \varepsilon \tag{10}$$

TABLE 1  
Comparison of time complexity of various PSHA algorithms

Algorithm	$N_m$	$N_r$	$N_\varepsilon$	$N_s$	$N_a$
Riemann sum	$N$	$N$	$N$	-	1
Conventional MC	1	1	1	$N$	1
IS MC	1	1	1	$N$	1
VEGAS AIS (this study)	1	1	1	$N$	$N$

$N_m$ , the number of magnitude grids;  $N_r$ , the number of distance grids;  $N_\varepsilon$ , the number of  $\varepsilon$  grids;  $N_s$ , the number of  $(m, r, \varepsilon)$  samples;  $N_a$ , the number of ground motion intensity of interest (e.g., 0.1 g)

133 , where  $\Delta m$ ,  $\Delta r$ , and  $\Delta \varepsilon$  are grid step size for Riemann sum,  $N_m$ ,  $N_r$ , and  $N_\varepsilon$  are the total number of grids, satisfying  $N_x \Delta x$   
134  $= x_{\max} - x_{\min}$ , where  $x = \{m, r, \varepsilon\}$ . Note that  $f_{M,R,\varepsilon}(m_i, r_j, \varepsilon_k) \Delta m \Delta r \Delta \varepsilon$  is equivalent to the probability at  $(m, r, \varepsilon)$  such that  
135  $|m - m_i| < \Delta m/2$ ,  $|r - r_j| < \Delta r/2$ , and  $|\varepsilon - \varepsilon_k| < \Delta \varepsilon/2$ . The accuracy of the Riemann summation depends on the grid size.  
136 Utilizing finer grids enhances the accuracy of the summation. However, the computation time is inversely proportional to the  
137 grid step size and thus proportional to the number of grids. For three-dimensional PSHA summation, the computation time  
138 scales with  $N_m \times N_r \times N_\varepsilon$  (Table 1). Notably, since the distance PDF ( $f_R(r)$ ) cannot be analytically determined in practice,  
139 integration often extends over latitude ( $\phi$ ), longitude ( $\psi$ ), and depth ( $z$ ), increasing the dimensions from three ( $m, r, \varepsilon$ ) to  
140 five ( $m, \phi, \psi, z, \varepsilon$ ). Therefore, Riemann summation for seismic hazard becomes even more computationally intensive, a  
141 phenomenon known as "the curse of dimensionality" due to the exponential increase in computation time with the number  
142 of dimensions (Novak and Ritter, 1997).

143 **Conventional Monte-Carlo (MC)** This method simulates many synthetic earthquake ground motions and calculates PSHA  
 144 by assessing the frequency with which ground motion intensities exceed a certain threshold. MC PSHA is computed as

$$\begin{aligned}\hat{\lambda}(X > a) &= \frac{\nu}{N} \sum_{i=1}^N I(X_i > a | M_i, R_i, \mathcal{E}_i) \\ &= \frac{1}{T} \sum_{i=1}^N I(X_i > a | M_i, R_i, \mathcal{E}_i)\end{aligned}\quad (11)$$

145 , where  $X_i$  denotes the simulated ground motions, and  $N$  is the total number of samples.  $M_i, R_i, \mathcal{E}_i$  are random samples from  
 146  $f_{M,R,\mathcal{E}}(m, r, \varepsilon)$ ,  $T$  represents the equivalent catalog duration equal to  $N/\nu$ . MC PSHA is unbiased because the expectation of  
 147  $\hat{\lambda}(X > a)$  is the same as  $\lambda$ :

$$\begin{aligned}E[\hat{\lambda}] &= E \left[ \frac{\nu}{N} \sum_{i=1}^N I(X_i > a | M_i, R_i, \mathcal{E}_i) \right] \\ &= \frac{\nu}{N} \sum_{i=1}^N E [I(X_i > a | M_i, R_i, \mathcal{E}_i)] \\ &= \frac{1}{N} \sum_{i=1}^N \nu \int I(X_i > a | m_i, r_i, \varepsilon_i) f_{M,R,\mathcal{E}}(m_i, r_i, \varepsilon_i) dm_i dr_i d\varepsilon_i \\ &= \frac{1}{N} \sum_{i=1}^N \lambda(X > a) \\ &= \lambda(X > a)\end{aligned}\quad (12)$$

148 The variance of  $\hat{\lambda}$ , the variability of each MC estimates, can be obtained as follows (Appendix A):

$$\text{VAR}[\hat{\lambda}] = \frac{\nu\lambda - \lambda^2}{N}\quad (13)$$

149 Note that  $\text{VAR}[\hat{\lambda}]$  is always positive since  $\nu \geq \lambda$ . Then, the standard deviation of the estimate,  $\hat{\lambda}$ , can be obtained as:

$$\sigma[\hat{\lambda}] = \sqrt{\frac{\nu\lambda - \lambda^2}{N}}\quad (14)$$

150 , which provides the absolute uncertainty about MC PSHA estimates. However, we do not want to fully rely on  $\sigma[\hat{\lambda}]$  to  
 151 compare different exceedance probabilities. For example, suppose we are interested in two different exceedance frequencies,  
 152  $\lambda_1 = 10^{-1}/\text{yr}$  and  $\lambda_2 = 10^{-2}/\text{yr}$  when  $\nu = 1/\text{yr}$  and  $N = 100$ . Then, the variance (uncertainty) of the two MC estimates,  $\hat{\lambda}_1$   
 153 and  $\hat{\lambda}_2$ , are  $\sigma_1 = 3 \times 10^{-2}$  and  $\sigma_2 \sim 1 \times 10^{-2}$ , respectively (Eq. (14)). Here, someone might argue that uncertainty of  $\hat{\lambda}_1$  MC  
 154 estimate is larger than that of  $\hat{\lambda}_2$  because  $\sigma_1$  is greater than  $\sigma_2$ . However, the uncertainty of  $\lambda_1$  should be considered smaller  
 155 than that of  $\lambda_2$  considering the target true value of each estimate. That is,  $0.1 \pm 0.03$  ( $\hat{\lambda}_1$ ) is better estimate than  $0.01 \pm 0.01$   
 156 ( $\hat{\lambda}_2$ ). Instead, it is better to use the coefficient of variation, COV, the standard deviation normalized by its mean



$$COV = \frac{\sigma[\hat{\lambda}]}{E[\hat{\lambda}]} = \sqrt{\frac{\nu - \lambda}{N\lambda}} \quad (15)$$

157 Smaller COV indicates that the PSHA estimate is more accurate relative to its actual value, e.g., 5% COV means that the  
 158 hazard estimates are within  $\pm 5\%$  of its true value with a 68% probability, assuming normally distributed MC estimates as per  
 159 the central limit theorem. COV is inversely proportional to the square root of the number of MC samples ( $N$ ), indicating that  
 160 increasing  $N$  naturally improves the estimate's accuracy.

161 The target exceedance frequency,  $\lambda$ , also affects the MC accuracy. The lower  $\lambda$  leads to poor accuracy with fixed  $N$  and  $\nu$ . It  
 162 is intuitively reasonable that sufficiently long earthquake catalog is required to accurately estimate the event from long return  
 163 period. It is also notable that the variance of MC estimate is proportional to the square root of earthquake occurrence rate,  
 164  $\nu$ , indicating that MC PSHA is more challenging task in region with higher earthquake activity. This can also be explained  
 165 intuitively. Regions with high seismic activity experience more earthquakes per year than less active areas, therefore, when  
 166 calculating the ground motion exceedance probability over the same period, more earthquakes need to be considered in these  
 167 active regions.

168 An advantage of MC PSHA is that the computational time is dependent on the number of samples  $N$ , circumventing the  
 169 inherent dimensionality problem in Riemann summation (Table 1). In other words, the computation time of MC simulation  
 170 is independent of the number of grids ( $N_m$ ,  $N_r$ , and  $N_\epsilon$  in Eq. (10)). Therefore, we could utilize fine joint probability mass  
 171 functions for more precise hazard estimation without an increase in computational burden. If closed-form probability density  
 172 functions are available, PSHA can be implemented without approximated discretization. Also, the computation time of MC is  
 173 independent of the number of ground motions of interest,  $N_a$  (Table 1). We can utilize the generated synthetic ground motion  
 174 catalog to tally the exceedance events for all the ground motions of interest, though we still should repeat the computation  
 175 along with the number of MC samples.

176 Conventional MC faces extreme computational challenges for low probabilities because the number of samples required  
 177 to achieve low COVs dramatically increases. We can compute the required number of samples by rearranging Eq. (15):

$$N = \frac{1}{(COV)^2} \times \frac{\nu - \lambda}{\lambda} \quad (16)$$

178 For small  $\lambda$ , i.e.,  $\nu \gg \lambda$ , the Eq. (16) can be approximated as :

$$N \sim \frac{1}{(COV)^2} \times \frac{\nu}{\lambda} \quad (17)$$

179 To demonstrate conventional MC's extreme computational demands, we can calculate  $N$  for a target annual exceedance  
 180 frequency,  $\lambda$ , of  $10^{-4}$  per year and a annual earthquake rate  $\nu = 1/\text{yr}$ , which is typical values in PSHA practice (Coppersmith  
 181 et al., 2014). From Eq. (17), we need  $N \sim 10^8$  to achieve  $\text{COV} = 1\%$ .

## 182 Importance Sampling (IS) Integration

183 IS is a generalization of MC theory. Consider a random variable  $x$  that follows a probability function,  $f_X(x)$ . The expected  
 184 value of a function  $u(x)$ , denoted as  $S$ , is defined by

$$S = \int u(x)f_X(x)dx \quad (18)$$

185 The MC estimate is

$$\hat{S} = \frac{1}{N} \sum_{i=1}^N u(X_i) \quad (19)$$

186 , where  $X_i$  is sampled from  $f_X(x)$ .

187 By introducing an arbitrary probability function,  $q_X(x)$ , Eq. (18) can be equivalently expressed as

$$S = \int u(x) \frac{f_X(x)}{q_X(x)} q_X(x) dx \quad (20)$$

188 We restrict the integration range in Eq. (20) where  $f_X(x) \neq 0$  because  $x$  such that  $f_X(x) = 0$  does not contribute to the inte-  
 189 gration. Then,  $q_X(x)$  can be any distribution with nonzero density in the integration range. Eq. (20) provides important  
 190 implications in MC estimation. We can get the solution to Eq. (18) by estimating the expected value of  $u(x)f_X(x)/q_X(x)$   
 191 where  $x$  follows the distribution  $q_X(x)$ . This approach is highly useful in numerical integration, especially when sampling  
 192 from  $f_X(x)$  is challenging or the population of  $f_X(x)$  is extremely low in the region of importance, e.g., the exceedance of  
 193 PGA greater than 1 g is mostly contributed by the ground motion samples from large magnitude ( $m$ ) earthquakes occurred  
 194 at close distance ( $r$ ) with large  $\epsilon$ , which all typically correspond to the low probability region.

195 The IS MC estimate is:

$$\hat{S} = \frac{1}{N} \sum_{i=1}^N u(X_i) \frac{f_X(X_i)}{q_X(X_i)} \quad (21)$$

196 , where  $X_i$ s are sampled from  $q_X(x)$ , the proposed (or new) IS sampling density function. If  $q_X(x)$  is equal to the original  
 197 distribution,  $f_X(x)$ , Eq. (21) simplifies to the conventional Monte-Carlo (Eq. (19)). The ratio  $f_X(x)/q_X(x)$ , known as the  
 198 importance weight ( $w_i$ ), adjusts for the change in sampling distribution.

199 **IS PSHA** Researchers have applied IS to PSHA using different sampling functions (Jayaram and Baker, 2010; Rahimi and  
 200 Mahsuli, 2019). For IS PSHA, Eq. (1) can be reformulated by introducing a new sampling joint density function,  $q_{M,R,\epsilon}(m, r, \epsilon)$

201 as follows

$$\lambda(X > a) = \nu \int_{\varepsilon_{\min}}^{\varepsilon_{\max}} \int_{r_{\min}}^{r_{\max}} \int_{m_{\min}}^{m_{\max}} I(X > a|m, r, \varepsilon) \frac{f_{M,R,\varepsilon}(m, r, \varepsilon)}{q_{M,R,\varepsilon}(m, r, \varepsilon)} q_{M,R,\varepsilon}(m, r, \varepsilon) dm dr d\varepsilon \quad (22)$$

202 The IS MC estimator of equation (22) is

$$\hat{\lambda}(X > a) = \frac{\nu}{N} \sum_{i=1}^N I(X_i > a|M_i, R_i, E_i) \frac{f_{M,R,\varepsilon}(M_i, R_i, E_i)}{q_{M,R,\varepsilon}(M_i, R_i, E_i)}$$

203 The mean of IS estimator  $\hat{\lambda}$  is also unbiased like conventional MC because

$$\begin{aligned} E_q[\hat{\lambda}] &= \frac{\nu}{N} \sum_{i=1}^N E_q \left[ I(X_i > a|M_i, R_i, \varepsilon_i) \frac{f_{M,R,\varepsilon}(M_i, R_i, \varepsilon_i)}{q_{M,R,\varepsilon}(M_i, R_i, \varepsilon_i)} \right] \\ &= \frac{1}{N} \sum_{i=1}^N \nu \int I(X > a|m, r, \varepsilon) \frac{f_{M,R,\varepsilon}(m, r, \varepsilon)}{q_{M,R,\varepsilon}(m, r, \varepsilon)} q_{M,R,\varepsilon}(m, r, \varepsilon) dm dr d\varepsilon \\ &= \frac{1}{N} \sum_{i=1}^N \lambda \\ &= \lambda \end{aligned} \quad (23)$$

204 Also, the variance of the IS PSHA estimates with respect to the true  $\lambda$  is given by

$$\text{VAR}[\hat{\lambda}] = \frac{1}{N} \left( \nu^2 E \left[ \left( I(X_i > a|M_i, R_i, \varepsilon_i) \frac{f_{M,R,\varepsilon}(M_i, R_i, \varepsilon_i)}{q_{M,R,\varepsilon}(M_i, R_i, \varepsilon_i)} \right)^2 \right] - \lambda^2 \right) \quad (24)$$

205 Hence, the COV of IS PSHA estimate can be expressed as

$$\text{COV} = \sqrt{\frac{\nu^2 E \left[ \left( I(X_i > a|M_i, R_i, \varepsilon_i) \frac{f_{M,R,\varepsilon}(M_i, R_i, \varepsilon_i)}{q_{M,R,\varepsilon}(M_i, R_i, \varepsilon_i)} \right)^2 \right] - \lambda^2}{N\lambda^2}} \quad (25)$$

## 206 **Equivalency of optimal IS density and hazard deaggregation**

207 From Eq. (24), we can specifically choose a new sampling density  $q_{M,R,\varepsilon}^*$  that makes  $\text{VAR}[\hat{\lambda}] = 0$ ,

$$q_{M,R,\varepsilon}^* = \nu \frac{I(X_i > a|M_i, R_i, E_i) f_{M,R,\varepsilon}(M_i, R_i, E_i)}{\lambda} \quad (26)$$

208 Using  $q_{M,R,\varepsilon}^*$ , we could compute the true hazard,  $\lambda$ , with only one MC sample because IS MC is unbiased (Eq. (23)) and  
209 the variance is zero. We call  $q^*$  the optimal IS density for PSHA calculation.

210 We note that Eq. (26) is exactly the same as Eq. (9), indicating that the optimal density,  $q^*$ , is identical to the hazard  
211 deaggregation. That is, if we find  $q^*$ , we are able to not only dramatically enhance the computational efficiency of seismic  
212 hazard estimation but also obtain the hazard deaggregation distributions as a by-product.

213 In fact, we can also see this profound relationship between the hazard and deaggregation estimates by rearranging Eq. (9):

$$\lambda = \nu \cdot \frac{I(X > a|m, r, \varepsilon)f_{M,R,\varepsilon}(m, r, \varepsilon)}{f(m, r, \varepsilon|X > a)}$$

214 Because this identity holds for any values of  $(m, r, \varepsilon)$ , we can obtain the hazard at ground motion level ,  $a$ , with any  $(m, r, \varepsilon)$   
215 triplet if we know  $f(m, r, \varepsilon|X > a)$ . The term  $f_{M,R,\varepsilon}(m, r, \varepsilon)/f(m, r, \varepsilon|X > a)$  can be interpreted as the importance weight of  
216 IS, and  $f(m, r, \varepsilon|X > a)$ , the hazard deaggregation, is the optimal density.

217 In existing numerical methods, the deaggregation can be more computationally expensive than the hazard. In Riemann  
218 sum, we must save each sum element in memory and allocate those elements into appropriate deaggregation bins. Also, in  
219 conventional MC, we should save the long synthetic ground motion catalog with the corresponding  $(m, r, \varepsilon)$  triplet to allocate  
220 those into the proper bins. These operations necessitate significant computational memory and time. Therefore, the property  
221 that the optimal density resembles the hazard deaggregation can be considered a huge benefit for hazard analysts. IS MC  
222 with different densities other than  $q^*$  could also improve the computational efficiency compared to the conventional MC;  
223 however, it cannot give any information on the hazard deaggregation (Rahimi and Mahsuli, 2019; Jayaram and Baker, 2010).

224 Though the use of optimal density is a major benefit in computation of seismic hazard and hazard deaggregation, obtaining  
225 it is not trivial because  $I(X_i > a|M_i, R_i, E_i)$  and  $\lambda$  in Eq. (26) are unknowns before calculation of PSHA. Thus, we propose a  
226 new PSHA computation method to find it.

## 227 ADAPTIVE IMPORTANCE SAMPLING PSHA

228 In “adaptive” importance sampling (AIS), we iteratively train the IS density to find the optimal one ( $q^*$ ) by exploring impor-  
229 tant regions to compute  $I(X_i > a|M_i, R_i, E_i)$  and  $\lambda$  with a reduced number of MC samples. AIS must balance different factors  
230 in determining how many samples should be used to train optimal IS density. Fewer MC samples can make AIS fail as they  
231 will not allow us to explore the important regions effectively. On the other hand, many samples can impose large compu-  
232 tational demands, even larger than those from conventional MC. In AIS, we must also consider computational costs are  
233 proportional to the number of iterations for convergence. Thus, it is important to select appropriate algorithms that converge  
234 fast.

235 **VEGAS Formulation for PSHA** VEGAS is a non-parametric adaptive importance sampling (AIS) algorithm that iteratively  
236 identifies the optimal proposal density,  $q^*$  (Lepage, 1978, 2021). The algorithm has been developed and widely used in com-

237 putational physics (Kersevan and Richter-Was, 2013; Alwall et al., 2014), and is currently applied to chemistry, astrophysics,  
 238 finance, and medical statistics (Campolieti and Makarov, 2007; Garberoglio and Harvey, 2011; Ray et al., 2011; Sanders,  
 239 2014).

240 The VEGAS algorithm is conceptually straightforward and is recognized for its rapid convergence, especially when the  
 241 random variables involved are independent (Lepage, 1978, 2021). In PSHA, the variable  $\varepsilon$  is always considered an independent  
 242 variable (Eq. (3)). Additionally, the variables  $m$  and  $r$  are also treated as independent under point source assumption (Eq. (4)).  
 243 Notice, however, that when the finite-fault rupture model, in which the rupture dimension changes with magnitude, is  
 244 adopted,  $m$  and  $r$  can be correlated, and the distribution of  $r$  is conditional on the magnitude  $m$ . Also, note that the estimate  
 245 remains unbiased in this case, because its calculation is still within the IS framework (Eq. (23)).

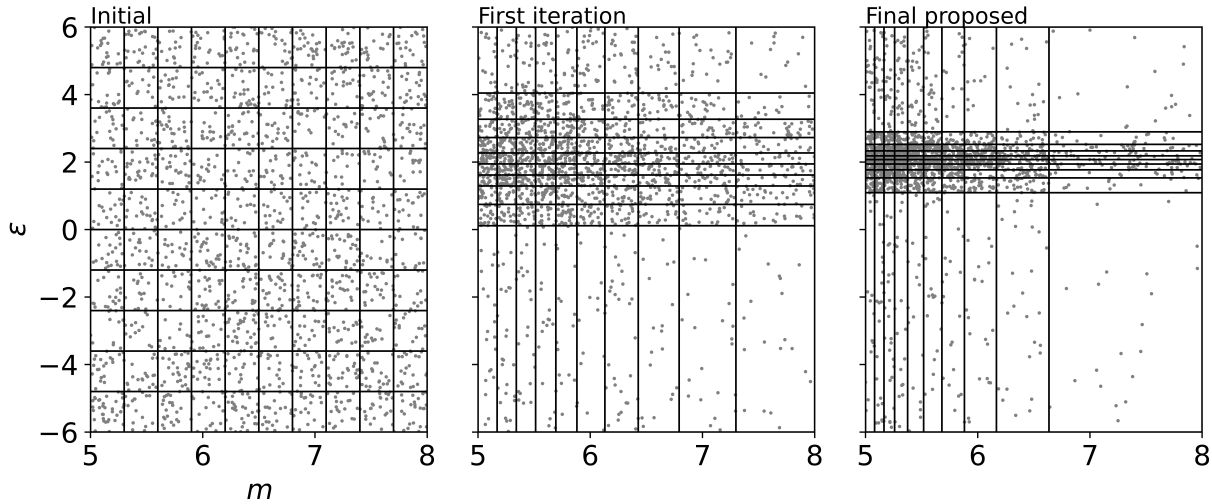
246 In three-dimensional integration, which is the case of PSHA, VEGAS employs  $N^3$  cuboids that are independently parti-  
 247 tioned. The probability assigned to each cuboid and the total number of partitioned cuboids ( $N^3$ ) are preserved across the  
 248 iteration steps. However, the IS sampling density changes because the algorithm updates the cuboid's size depending on  
 249 its contribution to the integration. If a cuboid's contribution is low, its size grows in the next step, lowering its probability  
 250 density. Conversely, when a cuboid's contribution is high, it shrinks in the following step, elevating the probability density.  
 251 Ideally, when every cuboid's contribution to the integration becomes identical, we find the proposed optimal density,  $q^*$ , and  
 252 the algorithm is terminated. The framework to find optimal IS densities for PSHA using VEGAS algorithm is explained with  
 253 a simple point seismic source example in the following paragraphs.

254 First, we adopt an independently distributed joint probability function as IS density. Note that we could also include  
 255 correlations in the IS density, but such approach would increase computational memory and time demands, e.g., the com-  
 256 putational complexity increases exponentially with each added dimension, i.e.,  $O(N^d)$  (Lepage, 1978). In contrast, when we  
 257 assume independence, the computational complexity grows linearly with the number of dimensions, i.e.,  $O(Nd)$ , making  
 258 multi-dimensional integration in PSHA exceptionally efficient. Thus,

$$q(m, r, \varepsilon) = q_M(m)q_R(r)q_\varepsilon(\varepsilon) \quad (27)$$

259 Then, the integration ranges for each variable— $m$ ,  $r$ , and  $\varepsilon$ —are divided into  $N$  grids with the same volume. This division  
 260 is designed to generate cuboids of constant probability:

$$\begin{aligned} M : m_{i-1} \leq m < m_i \quad (i = 0, 1, 2, \dots, N), \Delta m_i = m_i - m_{i-1} \\ R : r_{j-1} \leq r < r_j \quad (j = 0, 1, 2, \dots, N), \Delta r_j = r_j - r_{j-1} \\ \varepsilon : \varepsilon_{k-1} \leq \varepsilon < \varepsilon_k \quad (k = 0, 1, 2, \dots, N), \Delta \varepsilon_k = \varepsilon_k - \varepsilon_{k-1} \end{aligned} \quad (28)$$



**Figure 1.** VEGAS iterations of IS density in our AIS PSHA framework. Example for PSHA at  $PGA = 0.5$  g when the site is located 10 km away from a point seismic source ( $m_{\min} = 5.0$ ,  $m_{\max} = 8.0$ ,  $b$ -value = 1.0,  $\nu = 1.0/\text{yr}$ , GMM = [Sadigh et al. \(1997\)](#)). The total number of MC samples is 2,000, and the number of grids per axis is 10, constituting total 100 rectangles (cuboids in actual three dimensional PSHA integration). The vertical and horizontal black solid lines are boundaries of the  $m$  and  $\varepsilon$  rectangles, and gray dots are the MC samples. Starting from the initial same-size  $m$  and  $\varepsilon$  rectangles, their sizes are adjusted depending on the the contribution of each rectangle to the hazard. The final proposed structure of the rectangles gives a highly concentrated probability density at  $m \sim 5$  and  $\varepsilon \sim 2$ . Note that the hazard estimates using initialized density, density after the first iteration, and the final proposed density are 0.0417, 0.0413, and 0.0378, while the true solution is 0.0385

261 The number of grids,  $N$ , is chosen to be 50 as suggested by [Lepage \(1978\)](#) and because we consider this value makes grids  
 262 sufficiently fine to capture the actual distribution of the optimal density  $q^*$ . The probability for each cuboid from Eq. (28) is  
 263 set to  $1/N^3$  to ensure that the initial IS density function is uniformly distributed across the entire domain. As a result, the  
 264 initial probability density  $q^{(0)}$  for a specific cuboid is

$$q^{(0)}(m_{i-1} \leq m < m_i, r_{j-1} \leq r < r_j, \varepsilon_{k-1} \leq \varepsilon < \varepsilon_k) = q_M^{(0)}(m) q_R^{(0)}(r) q_\varepsilon^{(0)}(\varepsilon) = \frac{1}{N\Delta m_i} \frac{1}{N\Delta r_j} \frac{1}{N\Delta \varepsilon_k}$$

265 Figure 1 illustrates an example of initial partitioning when  $N=10$ . The gray dots in the figure shows an example of MC  
 266 samples from this initial  $q^0$ . Note that for easier visualization and understanding, Figure 1 illustrates an example in  $m$ - $\varepsilon$ , the  
 267 two-dimensional space, and not the three-dimensional  $m$ ,  $r$ , and  $\varepsilon$  space.

268 We update  $q$  using MC samples to ultimately make it converge to  $q^*$  (Eq. (26)). As mentioned earlier, the size of each cuboid  
 269 is subject to change while the probability of each cuboid remains constant, i.e., probability density changes. This adjustment  
 270 is facilitated through a "subdivision-and-restoration" process ([Lepage, 1978](#)). In this process,  $i$ th grid is subdivided into  $n_i$   
 271 sub-grids, with  $n_i$  being proportional to the  $i$ th grid's contribution to the overall integration, and restored to the original  
 272 number,  $N$ , by merging  $N_{\text{subgrid}}/N$  consecutive subgrids, where  $N_{\text{subgrid}}$  is the total number of sub-grids. Thus, the number

273 of subdivisions at  $i$ th grid is

$$n_i = N_{\text{subgrid}} \times \left( \frac{\bar{H}_i \Delta m_i}{\sum_i \bar{H}_i \Delta m_i} \right), \quad (i = 1, 2, \dots, N) \quad (29)$$

274

$$n_j = N_{\text{subgrid}} \times \left( \frac{\bar{H}_j \Delta r_j}{\sum_j \bar{H}_j \Delta r_j} \right), \quad (j = 1, 2, \dots, N) \quad (30)$$

275

$$n_k = N_{\text{subgrid}} \times \left( \frac{\bar{H}_k \Delta \varepsilon_k}{\sum_k \bar{H}_k \Delta \varepsilon_k} \right), \quad (k = 1, 2, \dots, N) \quad (31)$$

276  $N_{\text{subgrid}}$  should be sufficiently larger than  $N$  to iterate the IS density effectively, especially when the grids' contributions to  
 277 the hazard from the previous stage are highly heterogeneous. In this study, we chose  $N_{\text{subgrid}}$  to be 10,000, which is 200 times  
 278 greater than  $N$  ( $=50$ ). The second term of Eq. (29)-Eq. (31) represents the portion of each grid's contribution to the hazard.

279 According to Lepage (1978),  $\bar{H}_i$  in Eq. (29) is

$$\bar{H}_i = \sqrt{\sum_{m_{i-1} < M < m_i} \frac{H^2(M, R, \mathcal{E}; a)}{q_r(R)q_\varepsilon(\mathcal{E})}} \quad (32)$$

280 where

$$H(M, R, \mathcal{E}; a) = I(X > a | M, R, \mathcal{E}) f_{M,R,\mathcal{E}}(m, r, \varepsilon)$$

281 Note that  $H(M, R, \mathcal{E}; a)$  is the integrand of PSHA. Intuitively,  $\bar{H}_i$  can be considered as the marginalization of the overall  
 282 contribution of the magnitude dimension within the  $i$ th grid. Note that the summation of  $H^2$ 's in Eq. (32) is done over all the  
 283 samples which fall within  $m_{i-1}$  and  $m_i$  and the division by  $q_r q_\varepsilon$  can be considered an adjustment for unevenly distributed  $r$   
 284 and  $\varepsilon$  samples to purely capture the magnitude contribution.

285 Similarly,  $\bar{H}_j$  and  $\bar{H}_k$  in Eq. (30) and Eq. (31) can be calculated as

$$\bar{H}_j = \sqrt{\sum_{r_{i-1} < R < r_i} \frac{H^2(M, R, \mathcal{E}; a)}{q_m(M)q_\varepsilon(\mathcal{E})}}$$

286

$$\bar{H}_k = \sqrt{\sum_{\varepsilon_{i-1} < \mathcal{E} < \varepsilon_i} \frac{H^2(M, R, \mathcal{E}; a)}{q_r(R)q_m(M)}}$$

287 In the early iterations, researchers have noticed numerical instabilities due to the relatively poor information on the inte-  
 288 grand (Lepage, 1978, 2021). However, researchers have also found effective ways to mitigate it through smoothing (Eq. (33))

289 and damping (Eq. (34)).

$$d_i := \frac{1}{\sum d_i} \times \begin{cases} (7d_i + d_{i+1})/8 & \text{if } i = 0 \\ (d_{i-1} + 6d_i + d_{i+1})/8 & \text{if } i = 1, 2, \dots, N - 1 \\ (d_{i-1} + 7d_i)/8 & \text{if } i = N \end{cases} \quad (33)$$

290 , where  $d_i$  is  $\bar{H}_i \Delta x_i / \sum_i \bar{H}_i \Delta x_i$ , and  $x$  are  $m$ ,  $r$ , or  $\varepsilon$  in Eq. (29)-(31).

$$d_i := \left( \frac{1 - d_i}{\ln(1/d_i)} \right)^\alpha \quad (34)$$

291 , where  $\alpha$  is learning rate. We used  $\alpha$  as 1.0 as suggested by [Lepage \(2021\)](#). Through numerical experimenting, we observed  
292 this value allowed most of the PSHA integration to converge within three iterations (as described later) without causing any  
293 numerical instabilities.

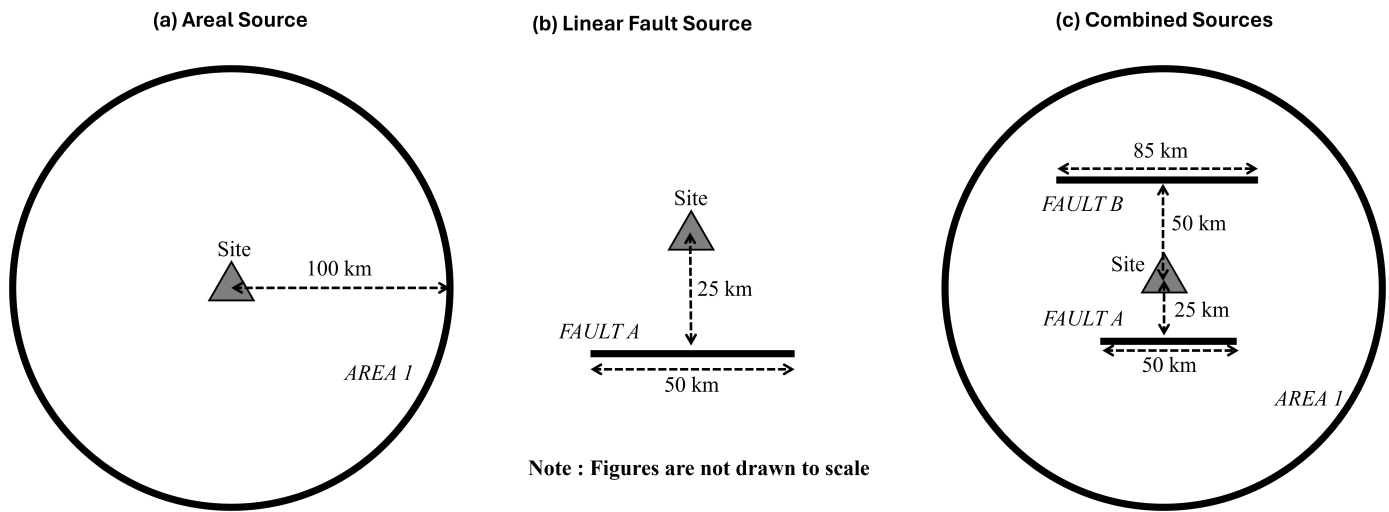
294 Finally, based on calculated  $n_i$ ,  $n_j$  and  $n_k$  (Eq. (29)-Eq. (31)), the number of grid is restored to the original size,  $N$ , by  
295 merging  $N_{\text{subgrid}}/N$  consecutive subgrids. The restored grid is the updated density,  $q^{(1)}$ . Figure 1 illustrates this iteration.  
296 The sizes of the grids are shrunk at  $5 < m < 6$  and  $\varepsilon \sim 2$ , indicating that the high contributions to the hazard on this range  
297 (Eq. (29)-Eq. (31)) in contrast to other less important regions, e.g., at  $m > 6.5$  and  $\varepsilon < 1$ . Note that the probability of each grid  
298 is preserved so that the number of MC samples inside each grid is almost the same regardless of the grid size.

299 The iterative process explained above is repeated until no further improvement is observed in the variance of the hazard  
300 estimator (Eq. (24)). An example of the grid structure of the final proposed IS density is shown in Figure 1. The final proposed  
301 density exhibits features which is expected based on intuition. First, the low contribution of  $\varepsilon$  less than 0 is understandable.  
302 This is because the simulated ground motion intensity, using [Sadigh et al. \(1997\)](#), never exceeds the target ground motion of  
303 0.5 g at any magnitude, even at  $m = 8.0$ , when  $\varepsilon$  is less than 0.1. In addition, the strong contribution of  $m \sim 5$  and  $\varepsilon \sim 2.2$  to  
304 the integration makes sense because the simulated ground motion intensity exceeds the target ground motion of 0.5 g when  
305  $m$  and  $\varepsilon$  reach 5 and 2.2, respectively. The decreasing trend of contribution beyond  $m \sim 5$  and  $\varepsilon \sim 2.2$  can be interpreted as the  
306 exponential decay in the probability of  $m$  and  $\varepsilon$ . This result for a simple PSHA example implies that the proposed algorithm  
307 can be utilized for more complex PSHA integration.

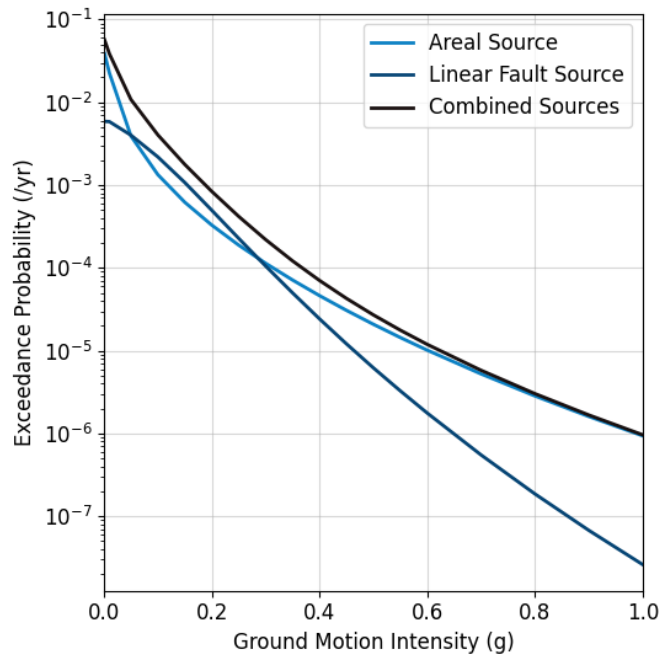
## 308 NUMERICAL EXAMPLES

309 We tested our proposed AIS PSHA on the comprehensive benchmark problem sets 1.11 and 2.1 from the PEER PSHA code  
310 verification project ([Hale et al., 2018](#)). For a thorough comparison, we performed numerical computations using four dif-  
311 ferent algorithms: Riemann sum, conventional Monte Carlo (MC), importance sampling (IS), and adaptive importance  
312 sampling (AIS) using the VEGAS algorithm. These computations were conducted across various seismic source settings:





**Figure 2.** Seismic source geometry for the three comprehensive numerical examples to test our AIS PSHA framework.



**Figure 3.** Benchmark PSHA curves for numerical examples 1 (areal source), 2 (linear fault source), and 3 (combined sources)

313 1) an areal source (Area1), 2) a fault source (FaultA), and 3) a combination of an areal and two fault sources (Area1, FaultA,  
 314 FaultB). These examples' geometry and seismic activity parameters are detailed in Figure 2 and Table 2. Figure 3 presents  
 315 the benchmark hazard curves for the examples. The ground motion model from Sadigh et al. (1997) was employed.

316 For Areal source 1, a circular-shaped areal source with a 100 km radius is considered (Figure 2). The earthquake activity  
 317 rate of the source,  $\nu(M > m_{\min})$ , is 0.0395/year with  $m_{\min}$  and  $m_{\max}$  of 5.0 and 6.5, respectively. The earthquake occurrence  
 318 model is doubly-truncated exponential with Gutenberg-Richter  $b$ -value of 0.9. Seismogenic depth is 5 to 10 km from the  
 319 surface, and a point source is assumed.

TABLE 2

Seismic source information for comprehensive numerical examples to test our AIS PSHA framework

	Area 1	Fault A	Fault B
Source Type	Area	Vertical Fault	Vertical Fault
Source Geometry	Circle (R=100 km)	Line (L = 50 km)	Line (L = 85 km)
Earthquake Occurrence Model	Exponential	Characteristic	Characteristic
b-value	0.9	0.9	0.9
$M_{\min}$	5.0	5.0	5.0
$M_{\max}$	6.5	6.75	7.0
$M_{char}$	-	6.5	6.75
$\nu(M > M_{\min})$ (yr <sup>-1</sup> )	0.0395	-	-
slip rate (mm/yr)	-	1	2
Seismogenic depth (km)	5-10	0-12	0-12
Rupture type	Point	Floating rectangular rupture	Floating rectangular rupture

320 Faults A and B are vertical fault sources (strike = 90°E, dip = 90°) with lengths of 50 and 85 km, respectively. They extend  
321 from the surface to 12 km depth. The earthquake activity models are chosen to be characteristic (Youngs and Coppersmith,  
322 1985) with the  $b$ -value of 0.9, and  $m_{\min}$  is set to be 5.0 for both fault sources. For Fault A, slip rate,  $M_{\max}$ ,  $M_{char}$  are 1 mm/year,  
323 6.75, and 6.5; and for fault B, they are 2 mm/year, 7.0, and 6.75. In fault sources, ruptures are assumed to be floating inside  
324 the fault with the rupture dimensions following

$$\log_{10}(A) = M - 4 \quad (35)$$

$$\log_{10}(W) = 0.5M - 2.15 \quad (36)$$

$$\log_{10}(L) = 0.5M - 1.85 \quad (37)$$

325 , where  $A$  is rupture area in km<sup>2</sup>,  $W$  is rupture width in km,  $L$  is rupture length in km, and  $M$  is earthquake magnitude. The  
326 probability of exceedance at eighteen ground motions (PGA), 0.001, 0.01, 0.05, 0.1, 0.15, 0.2, 0.3, 0.35, 0.4, 0.45, 0.5, 0.55, 0.6,  
327 0.7, 0.8, 0.9, and 1.0 g, are estimated.

328 For Riemann summation, we followed the calculation procedure specified in the PEER PSHA code verification project  
329 (Hale et al., 2018). Thus, we used magnitude and source-to-site distance step sizes of 0.01 and 0.1 km, respectively. Note  
330 that the number of grids of the source-to-site distance distribution ( $\sim 950$ ) is significantly less than that of source location  
331 probability, the uniformly distributed probability of event location across the seismic source region, with the spatial and  
332 depth grid spacing of 0.5 km and 1 km ( $\frac{\pi \times 100^2}{0.5 \times 0.5} \times 6 \approx 750,000$ ) suggested by Hale et al. (2018). This makes the Riemann sum  
333 computation  $\sim 800$  times more efficient. For the ground motion random variable ( $\epsilon$ ), the grid step size was set to 0.01, with  
334 minimum and maximum values of -6 and 6, respectively, beyond which the PSHA curve shows negligible change even at an  
335 annual exceedance probability of  $10^{-8}$  (Bommer and Abrahamson, 2006).

336 In the conventional MC approach, we generate a ground motion catalog based on the probability distributions of  $m$ ,  $r$ ,  
337 and  $\epsilon$ . From this catalog, we calculate the annual exceedance frequency by counting the instances where the ground motion

338 exceeds the specified threshold, as outlined in Eq. (11). This frequency is then converted to an annual exceedance probability  
 339 using Eq. (6).

340 For IS, we adopt uniform IS densities across the entire integration range for magnitude, distance, and  $\varepsilon$  following Jayaram  
 341 and Baker (2010). This choice is made due to the absence of prior information regarding which integration range significantly  
 342 influence the hazard calculation.

343 For our AIS PSHA with the VEGAS algorithm, we set the initial number of grids for  $m$ ,  $r$ , and  $\varepsilon$  to 50 (Eq. (28)). The total  
 344 number of subgrids,  $N_{\text{subgrid}}$ , is chosen to be 10,000. The learning rate,  $\alpha$ , is fixed at 1.0. Algorithm 1 presents the pseudocode.  
 345 From lines 1 to 5, the generation index  $t$  is set to be zero, and the sampling density  $q$  is set to be uniform. The main algorithm  
 346 loop is from lines 6 to 24. The loop is continued until there is no improvement in the coefficient of variation or there is no  
 347 previous generation (line 6). At line 7,  $X_m, X_r, X_\varepsilon$  are sampled from the distribution  $q^{(t)}$ . Then, the probability  $\hat{\lambda}$  is estimated  
 348 in lines 8-10, and its coefficient of variation is also calculated in line 11. From lines 12 to 23 the sampling function  $q^{(t)}$  is  
 349 updated. The steps are repeated over  $m, r$ , and  $\varepsilon$  (lines 12-15). From lines 16 to 20, the contribution of each grid is calculated,  
 350 at line 21, it is smoothed and dampened, and  $q^{(t)}$  is updated by subdivision depending on  $d_i$  and restore the number of grids to  
 351 the original number,  $N$ , in line 22. Updated  $q^{(t+1)}$  is obtained by multiplying  $q_m^{(t+1)}, q_r^{(t+1)}$ , and  $q_\varepsilon^{(t+1)}$  in line 23, and the while  
 352 loop is ended by increase the generation index  $t$ . When main algorithm loop is terminated, it returns the hazard estimate  $\hat{\lambda}$   
 353 and proposed optimal density  $q$  in line 25.

354 We assessed the accuracy of conventional MC, IS, and AIS probability estimates through the standard deviation of the  
 355 relative error with respect to the benchmark curve, calculated using the following formula:

$$\sigma = \sqrt{\frac{1}{N} \sum_{n=1}^N e_n^2 - \left( \frac{1}{N} \sum_{n=1}^N e_n \right)^2}$$

356 , where  $e_n$  is the relative error of the  $n$ th MC estimate, and  $N$  is the total number of MC exceedance probability estimates.  
 357 Since MC, IS MC, and VEGAS AIS estimates are unbiased (Eq. (12), Eq. (23)), the sum of  $e_n$  when  $N \rightarrow \infty$  is theoretically  
 358 zero. The relative error,  $e_n$ , is defined as:

$$e_n = \frac{p_n - p_{ref}}{p_{ref}}$$

359 , where  $p_n$  is the exceedance probability estimated at  $n$ th MC estimate, and  $p_{ref}$  is the benchmark probability.

360 The analyses were conducted in a Python 3.11 on an Intel Core i7-13700 2100 MHz processor with 64GB RAM.

### 361 **Example 1 : Areal source**

362 We consider the circular area source with the site at the circle's center in Figure 2. For this example, we show different compu-  
 363 tational performances to conduct PSHA in Figure 4. Our numerical experiments show that computational times and standard

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**Algorithm 1** VEGAS adaptive importance sampling PSHA pseudocode

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**[Parameters and Functions]**

$a$	Ground motion intensity of interest (e.g., $a = 0.1$ g)
$N$	Number of grids (e.g., $N = 50$ )
$N_{\text{subgrid}}$	Number of sub-grids (e.g., $N_{\text{subgrid}} = 10,000$ )
$N_s$	Number of MC samples (e.g., $N_s = 1,000$ )
$\vec{X}_m$	Magnitude sample vector ( $\vec{X}_{m,i} = i$ th element of $\vec{X}_m$ )
$\vec{X}_r$	Distance sample vector
$\vec{X}_\varepsilon$	Ground motion random variable sample vector
$\vec{X}$	$[\vec{X}_m; \vec{X}_r; \vec{X}_\varepsilon]$
$G(\cdot)$	Ground Motion Model
$I(\cdot)$	Indicator function
$f_X(\cdot)$	Original sampling distribution
$q^{(t)}(\cdot)$	Proposed sampling distribution at $t$ th iteration step
$\hat{\lambda}^{(t)}$	Hazard estimate at $t$ th iteration step
$\text{COV}^{(t)}$	COV of the hazard estimate at $t$ th iteration step
$\epsilon$	pre-defined iteration stopping criteria (e.g. COV of 0.02)
$i$	Index for partitioned grids ( $i = 1, 2, \dots, N$ )
$j$	Index for samples ( $j = 1, 2, \dots, N_s$ )

**[Algorithm]**

```
1:  $t = 0$ 
2:  $q_m^{(t)} = \text{U}([m_{\min}, m_{\max}])$ 
3:  $q_r^{(t)} = \text{U}([r_{\min}, r_{\max}])$ 
4:  $q_\varepsilon^{(t)} = \text{U}([\varepsilon_{\min}, \varepsilon_{\max}])$ 
5:  $q^{(t)} = q_m^{(t)} q_r^{(t)} q_\varepsilon^{(t)}$ 

6: while  $\text{COV}^{(t-1)} < \epsilon$  or  $t < 2$  :
7:    $\vec{X} = \{X_m, X_r, X_\varepsilon\}^{N_s}$ , ( $X_m \sim q_m^{(t)}$ ,  $X_r \sim q_r^{(t)}$ ,  $X_\varepsilon \sim q_\varepsilon^{(t)}$ )
8:    $\vec{H}^{(t)} = \{I(G(x) > a|x)f_X(x) \mid x \in \vec{X}\}$ 
9:    $\vec{\lambda}^{(t)} = \vec{H}^{(t)}/q^{(t)}(\vec{X})$ 
10:   $\hat{\lambda}^{(t)} = \sum_1^{N_s} \lambda_i^{(t)}/N_s$ 
11:   $\text{COV}^{(t)} \leftarrow \text{Eq. (25)}$ 
12:  for  $u$  in  $[m, r, \varepsilon]$  :
13:    if  $u = m$  :  $(v, w) \leftarrow (r, \varepsilon)$ 
14:    if  $u = r$  :  $(v, w) \leftarrow (\varepsilon, m)$ 
15:    if  $u = \varepsilon$  :  $(v, w) \leftarrow (m, r)$ 
16:    for  $i$  in  $\{1, 2, \dots, N\}$  :
17:      for  $j$  in  $\{1, 2, \dots, N_s\}$  :
18:        if  $u_{i-1} < X_{u,j} < u_i$  :  $\bar{H}_i \leftarrow \bar{H}_i + (H_j^2 / (q_v(X_{v,j}) \cdot q_w(X_{w,j})))$ 
19:         $\bar{H}_i \leftarrow \sqrt{\bar{H}_i}$ 
20:         $d_i \leftarrow \left( \frac{\bar{H}_i \Delta x_i}{\sum_i \bar{H}_i \Delta x_i} \right)$ 
21:         $d_i \leftarrow$  smoothed, dampened  $d_i$  (Eq. (33), (34))
22:         $q_u^{(t+1)} \leftarrow$  subdivision and restoration (Eq. (29)-(31))
23:         $q^{(t+1)} \leftarrow q_m^{(t+1)} q_r^{(t+1)} q_\varepsilon^{(t+1)}$ 
24:         $t \leftarrow t + 1$ 
25:  return  $\hat{\lambda}^{(t-1)}$ ,  $q^{(t-1)}$ 
```

---

364 deviations (“accuracy”) are linearly correlated in logarithmic scale, in agreement with the theory, because the required num-  
365 ber of MC samples (linearly proportional to the computational time) is inversely proportional to the square of the standard  
366 deviation (Eq. (16)).

367 While the computation time and standard deviation vary depending on the target ground motion of interest, our results  
368 show our AIS PSHA generally outperforms the other numerical techniques. AIS PSHA becomes increasingly efficient for  
369 larger ground motions. At a low target ground motion of 0.05g, the computation time to achieve a 2 % standard deviation  
370 is 0.01 seconds for AIS, while it takes 0.05 and 0.09 seconds for conventional MC and IS estimates, respectively, indicating  
371 AIS is 4.8 and 8.5 times faster. However, computational efficiency becomes extreme in high-ground motions. At 1.0 g, the  
372 computation time to achieve 2 % standard deviation is estimated to be 0.02 seconds for AIS, while it takes 127 and 1.4 seconds  
373 for conventional MC and IS estimates. In this case, AIS is 7,800 and 70 times faster than conventional MC and IS, respectively.  
374 Also, note that for 2 % standard deviation case, AIS is  $> 10^5$  faster than Riemann sum. We also note that AIS outperforms IS  
375 in all ground motion ranges by a factor of 8 to 70 to achieve a 2 % standard deviation. However, this is not always guaranteed  
376 because AIS takes  $t$  times more computation time than IS with the same  $N$  due to the  $t$  iterations to find the optimal IS  
377 density. Thus, our findings imply AIS PSHA with the VEGAS algorithm can find (close to) optimal IS density quickly.

378 It is also noteworthy that with fixed number of samples (i.e., constant computational burden), AIS’s estimation accuracy is  
379 quite similar across different target ground motions, while conventional MC’s accuracy decreases sharply for higher ground  
380 motions (Figure 5). In fact, if the hazard curve is exponential (Marzocchi and Jordan, 2017), we can show conventional MC  
381 decreases its efficiency also exponentially (Eq. (15)). For example, the error in the hazard curve for 1.0g will increase to 100%  
382 if it initially was 1 % for 0.001g with exceedance frequency of 0.9/year (when  $\nu = 1/\text{year}$ ). From our numerical experiments,  
383 we found errors grow from 0 to 5.60 % with  $N_{\text{convMC}} = 10^7$  MC samples for these ground motions. In contrast, for a similar  
384 computation time, we found AIS had errors ranging from 0.13 to 0.64 %, for this quite different ground motion levels (Figure  
385 5).

386 This finding is key for PSHA as the computational bottleneck is at the highest ground motion intensity. Using conventional  
387 MC PSHA, the hazard analyst has no choice but to largely increase the number of samples to estimate hazard accurately  
388 at high ground motions even though such a large number would not be necessary for low ground motions. This makes  
389 the conventional MC highly inefficient, and consequently, the efficiency of the conventional MC for lower ground motions  
390 cannot be considered a real advantage for PSHA. AIS PSHA overcomes this problem by adopting different optimal densities  
391 at different ground motion intensities, making the computational burden to achieve the similar estimation error almost flat  
392 for any ground motion intensity, as shown in Figure S1. In terms of the accuracy of AIS PSHA, the PEER PSHA verification  
393 project suggests a strict acceptable error range of 5% for reliable PSHA computation codes. AIS PSHA achieves it with only  
394  $N \sim 10,000$  per ground motion (Figure S1).

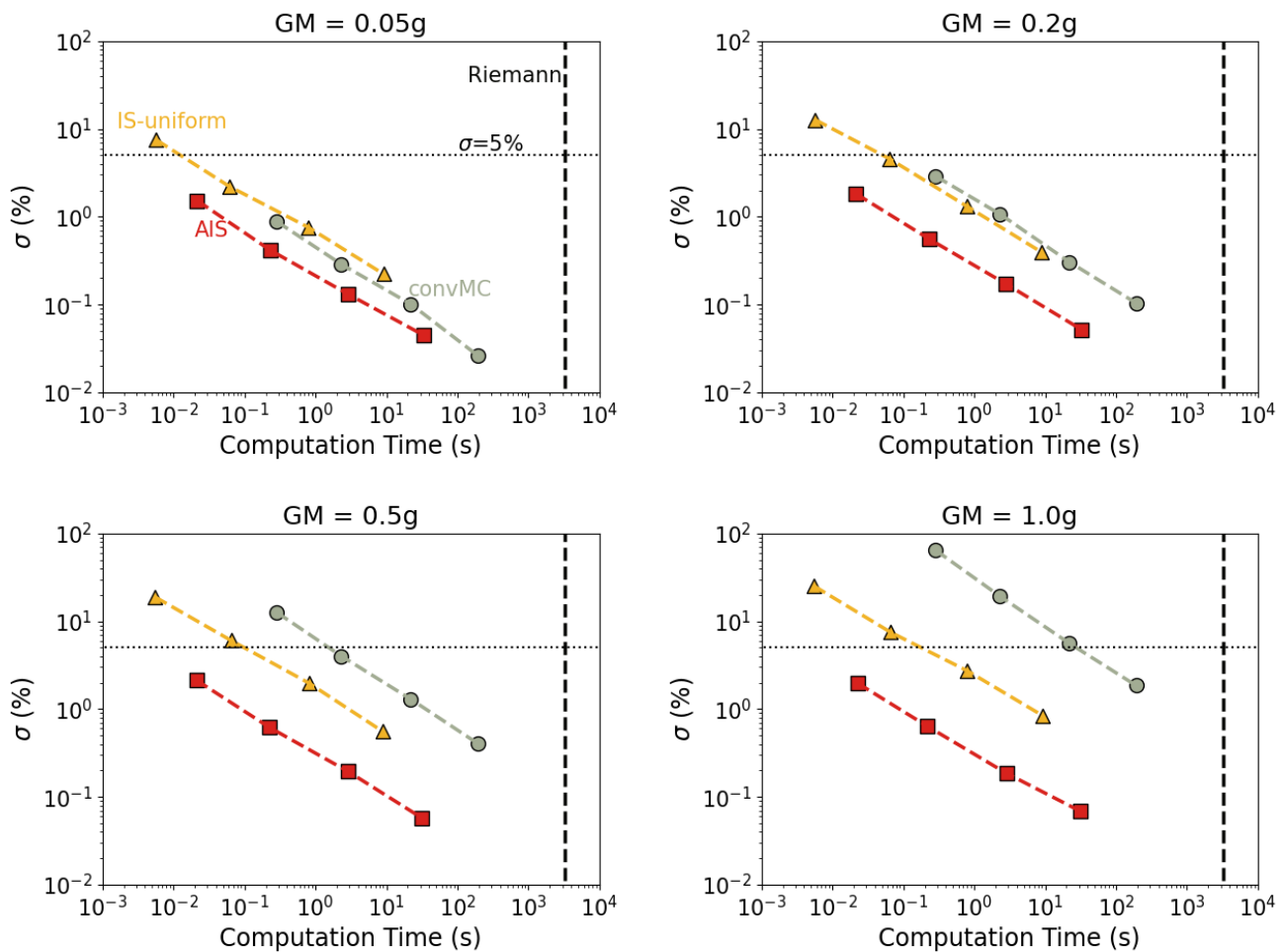
395 Another key advantage of our proposed AIS PSHA is the co-production of deaggregation curves at no extra computational  
396 cost. From Eq. (26), optimal IS densities are theoretically equivalent to hazard deaggregation distributions. We show AIS can  
397 find close-to-optimal IS densities and thus closely match hazard deaggregation curves. We found the benchmark marginal  
398 distributions of hazard deaggregation obtained from Riemann sum closely match the iterated IS density from AIS PSHA  
399 (Figure 6). We used the Kolmogorov-Smirnov (K-S)  $D$  statistic (Kolmogorov, 1933) to quantify their similarities. K-S  $D$  statistic  
400 measures the maximum difference between two cumulative distribution functions (CDF). If two CDF are identical,  $D$  is zero,  
401 and its maximum possible value is one.  $D$  close to zero indicates that two probability distributions are similar. We calculated  
402  $D$  for  $m$ ,  $r$ , and  $\varepsilon$  at all the ground motion intensities (Figure S2). We found the maximum  $D$  values for  $m$ ,  $r$ ,  $\varepsilon$  were 0.032,  
403 0.113, and 0.092, respectively, and the minimum values were 0.019, 0.026, and 0.017, indicating a strong resemblance between  
404 the two distributions.

405 We also estimated the differences of mean values from proposed  $q^*$  and Riemann sum hazard deaggregation (Figure S2).  
406 Note that the deaggregation distributions and optimal IS densities vary for different ground motion levels. Thus, their mean  
407 values also vary. We found the maximum relative differences were 2.5 %, 22.6 %, and 18.5 % for  $m$ ,  $r$ ,  $\varepsilon$ , and the minimum  
408 differences were 1.7 %, 1.0 %, and 4.3 %, respectively. Also, the maximum absolute differences were 0.14, 3.9 km, and 0.17  
409 for  $m$ ,  $r$ ,  $\varepsilon$ , and the minimum differences were 0.09, 0.7 km, and 0.0009. The maximum relative difference in distance ( $r$ )  
410 appears for deaggregation distributions at 0.25 g, where the mean distance obtained from hazard deaggregation is 14.3 km  
411 and that from  $q^*$  is 17.5 km. Given that we are typically interested in distance ranges on the order of tens of kilometers (e.g.,  
412 0-15 km, 15-25 km, 25-50 km, etc.) rather than a single value (U. S. Nuclear Regulatory Commission, 2007), this difference  
413 is not crucial in determining the controlling earthquake for critical infrastructures. In addition, IS densities still show small  
414 K-S  $D$  statistics and their mode almost matches each other, as shown in Figure 6.

## 415 **Example 2 : Fault source**

416 We consider a 50 km-length vertical fault 25 km away from the site (Figure 2). We adopted a finite-dimension rupture model,  
417 which results in a distance distribution dependent on magnitude. We found this dependency can diminish the performance  
418 of VEGAS AIS because the VEGAS algorithm assumes the independently distributed random variables. For example, for a  
419 ground motion intensity of 0.05 g, AIS is slower than conventional MC (Figure 7), while AIS outperformed conventional  
420 MC at the same ground motion intensity when point source assumption was made, as seen earlier (Figure 4). However, as  
421 the ground motion intensity increases, the computational gap between the two methods becomes smaller rapidly and closes  
422 at 0.2g. For higher ground motions, AIS outperforms conventional MC. At 1.0 g, for a 5% standard deviation, AIS, IS, and  
423 conventional MC take 0.08, 3.7, and 166 seconds, respectively, i.e., AIS MC is 48 and 2,162 times faster.

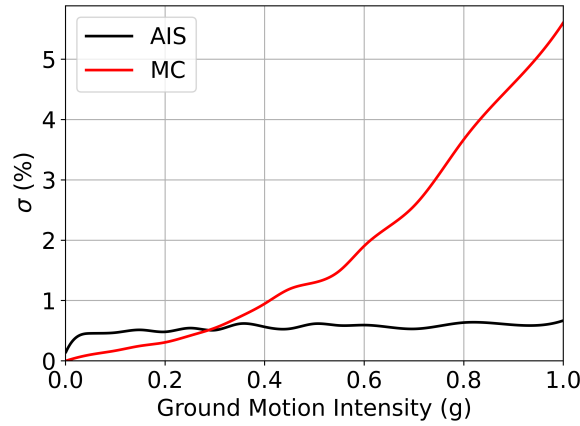
424 We also present the accuracy of AIS PSHA in comparison to the true hazard calculated by Riemann summation (Figure  
425 S3). We observed AIS PSHA estimates approximate the true hazard curve within an acceptable range (5 % error) when  $N$



**Figure 4.** Areal Source Example: Standard deviation of the conventional MC (red), IS MC (blue), AIS MC (black) estimates as functions of computation time at ground motion of 0.05 g, 0.2 g, 0.5 g, and 1.0 g. The computation time required for Riemann summation is presented as green vertical dashed line. At low ground motion, conventional MC outperforms IS and AIS MC, however, as the target ground motion increases, the performance of AIS make a dramatic improvement in terms of both accuracy and computational cost.

426 is greater than  $\sim 50,000$ . Also, note that though we assumed the independently distributed optimal density, AIS PSHA still  
 427 gives an unbiased hazard estimate that will converge to the true hazard with a sufficiently large number of samples due to  
 428 the nature of the importance sampling (Figure S3 (d)).

429 We also compared hazard deaggregation and the iterated IS density (Figure 8) and showed they closely match each other  
 430 even though the distance distribution depends on the magnitude in this case. It is also noteworthy the iterated IS density can  
 431 even reproduce complex densities with discontinuities like the large jump within magnitude distribution (for  $m = 6.25$ ) due to  
 432 the use of a characteristic earthquake occurrence model. The K-S D statistic and mean difference of the two distributions are  
 433 also presented in Figure S4. The maximum values of D in  $m$ ,  $r$ , and  $\varepsilon$  are 0.30, 0.68, and 0.13, respectively, and the minimum  
 434 values are 0.02, 0.31, and 0.04. We found that the largest discrepancies occur in the magnitude distribution, but errors can be  
 435 considered negligible as the mean magnitude difference is within 6 % error. We also note considerable discrepancies in the  
 436 distance distribution shape (see K-S D statistics of Figure S4) as curves with concentrated probabilities in narrow ranges are



**Figure 5.** Standard deviation of MC estimates with the ground motion for conventional (red) MC ( $N_{\text{convMC}} = 10^7$ ) and AIS (black) with the similar computation time. The standard deviation of error exponentially increase with the ground motion in conventional MC, while that of AIS remain constant.

437 harder to estimate for AIS (Figure 8). However, the mean distance is still within a 5 % error range. The  $\varepsilon$  is generally in good  
 438 agreement across all ground motion intensities. We observe higher differences at lower ground motion because the mean  $\varepsilon$   
 439 is close to zero. For example, the calculated mean  $\varepsilon$  at ground motion intensity of 0.01 g is 0.022 and 0.007, which is not a  
 440 large difference in practice.

441 The maximum relative differences in mean values were found to be 6.2 %, 5.3 %, and 129 % for  $m$ ,  $r$ ,  $\varepsilon$ , and the minimum  
 442 differences were 0.006 %, 3.9 %, and 2.2 %, respectively. Note that the 129 % of  $\varepsilon$  case is corresponding the case where the  
 443 mean  $\varepsilon$  is close to zero. The relative difference appears to be slightly higher than previous areal source example, however, the  
 444 absolute difference is still remain to be significantly small, the maximum absolute differences were 0.36, 1.3 km, and 0.20 for  
 445  $m$ ,  $r$ ,  $\varepsilon$ , and the minimum differences were 0.0004, 1.0 km, and 0.07.

### 446 Example 3 : Combined sources

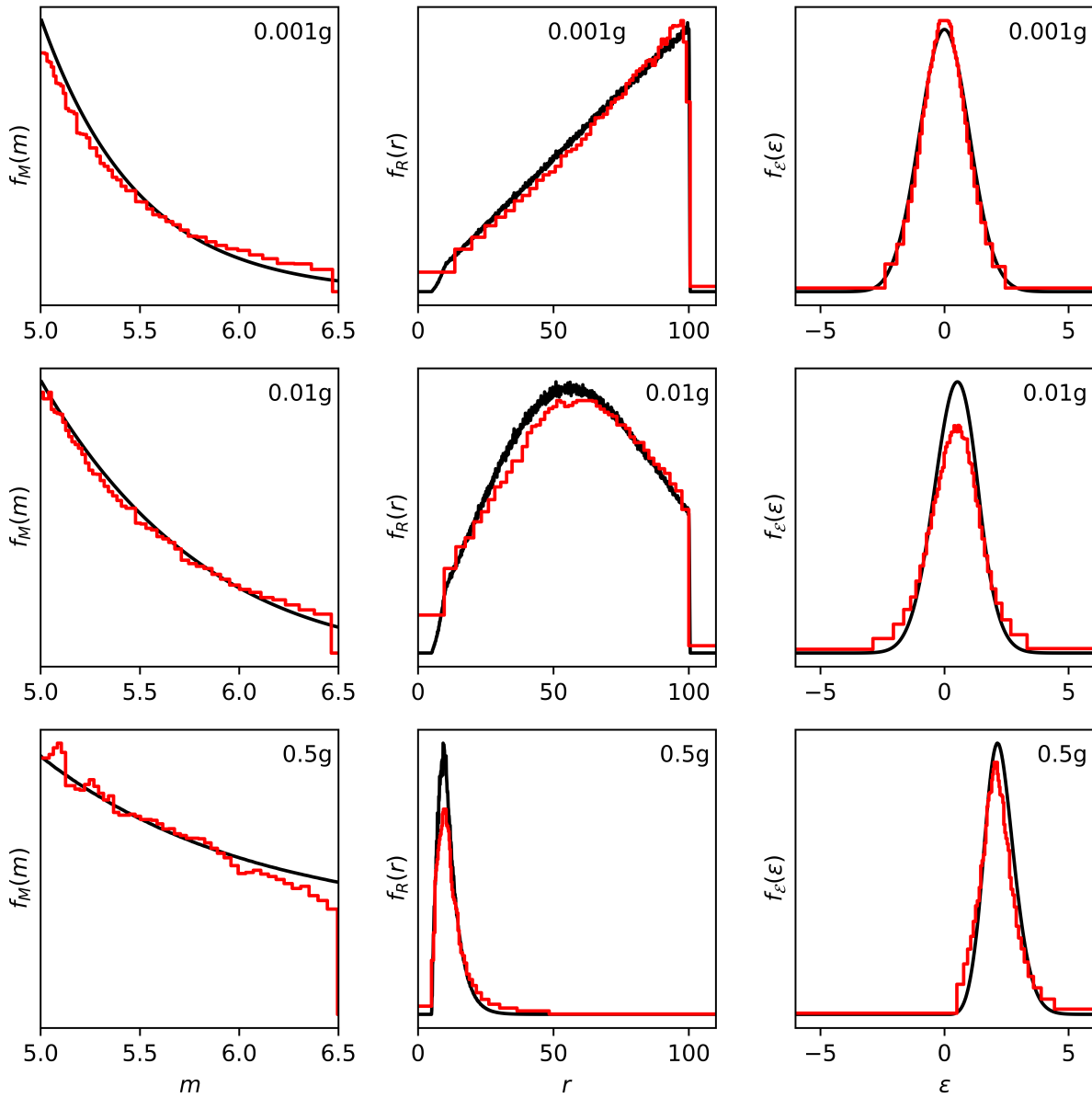
447 In PSHA, we often have multiple seismic sources. We consider one areal and two fault sources around the site to represent this  
 448 case. This application posits a different mathematical problem than the previous two examples because we must introduce  
 449 an additional variable to formulate AIS.

450 First, the probability of earthquake occurrence at  $i$ th seismic source can be defined as:

$$P(S = i) = \frac{\nu_i}{\sum_{i=1}^{N_s} \nu_j}$$

451 , where  $N_s$  is the number of seismic sources ( $N_s = 3$  in this example), and  $\nu_i$  is the annual earthquake occurrence rate of  $i$ th  
 452 seismic source. Because the discrete random variables cannot be used in AIS, we define a continuous random variable and  
 453 its corresponding probability density function as

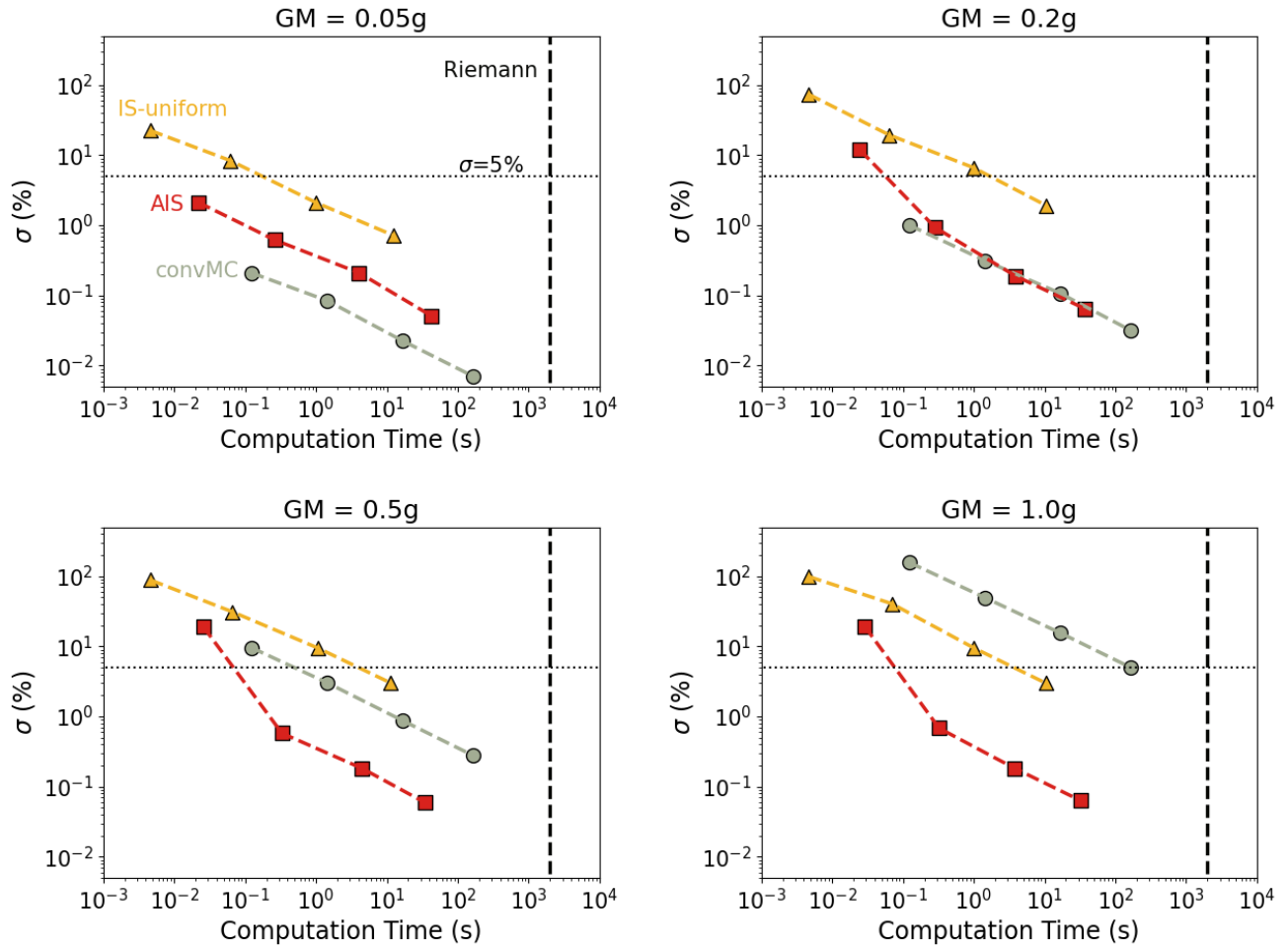




**Figure 6.** Areal Source Example: The convergence of  $m$ ,  $r$ , and  $\varepsilon$  iterated IS densities density (red) derived in AIS algorithm ( $N = 100,000$ ) to marginal distributions of hazard deaggregation (black) at ground motion intensities of 0.001 g, 0.01 g, and 0.5 g.

$$f_S(s) = \begin{cases} 0 & s < 0 \\ P(S=i) & \text{if } i-1 \leq s < i \text{ (} i=1, 2, \dots, N_s \text{)}, \\ 0 & N_s \leq s \end{cases} \quad (38)$$

454  $f_S(s)$  is a piece-wise constant function where the heights are proportional to the corresponding sources' earthquake  
 455 occurrence rates. We introduce  $f_S(s)$  into the PSHA integration and obtain

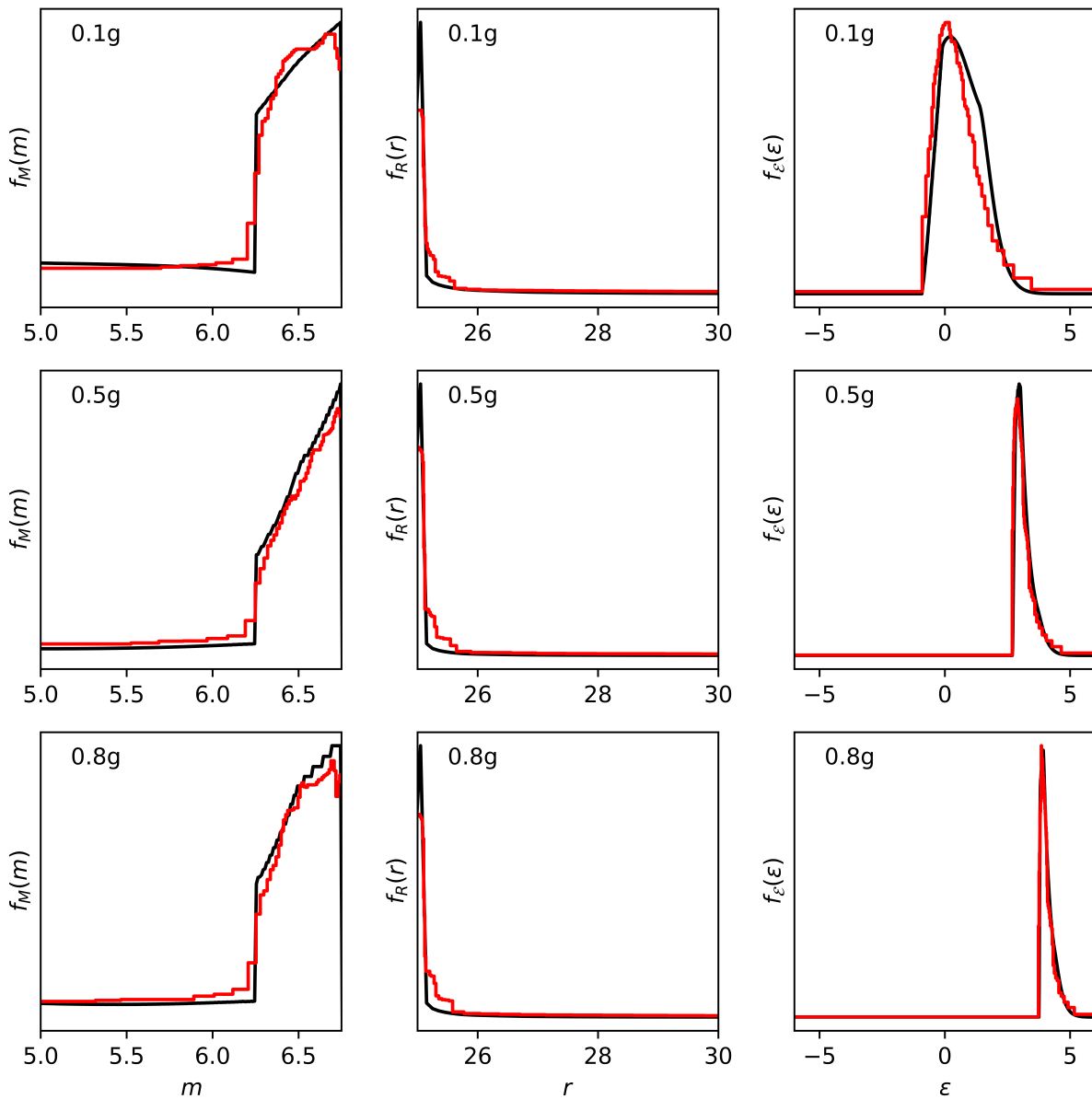


**Figure 7.** Fault Source Example: Standard deviation of the conventional MC (red), IS MC (blue), AIS MC (black) estimates as functions of computation time at ground motion of 0.05 g, 0.2 g, 0.5 g, and 1.0 g.

$$\Lambda(X > a) = \left( \sum_i^{N_{src}} \nu_i \right) \int_S \int_\epsilon \int_R \int_M I(x > a | s, m, r, \epsilon) f_{M,R,\epsilon}(m, r, \epsilon | s) f_S(s) dm dr d\epsilon ds \quad (39)$$

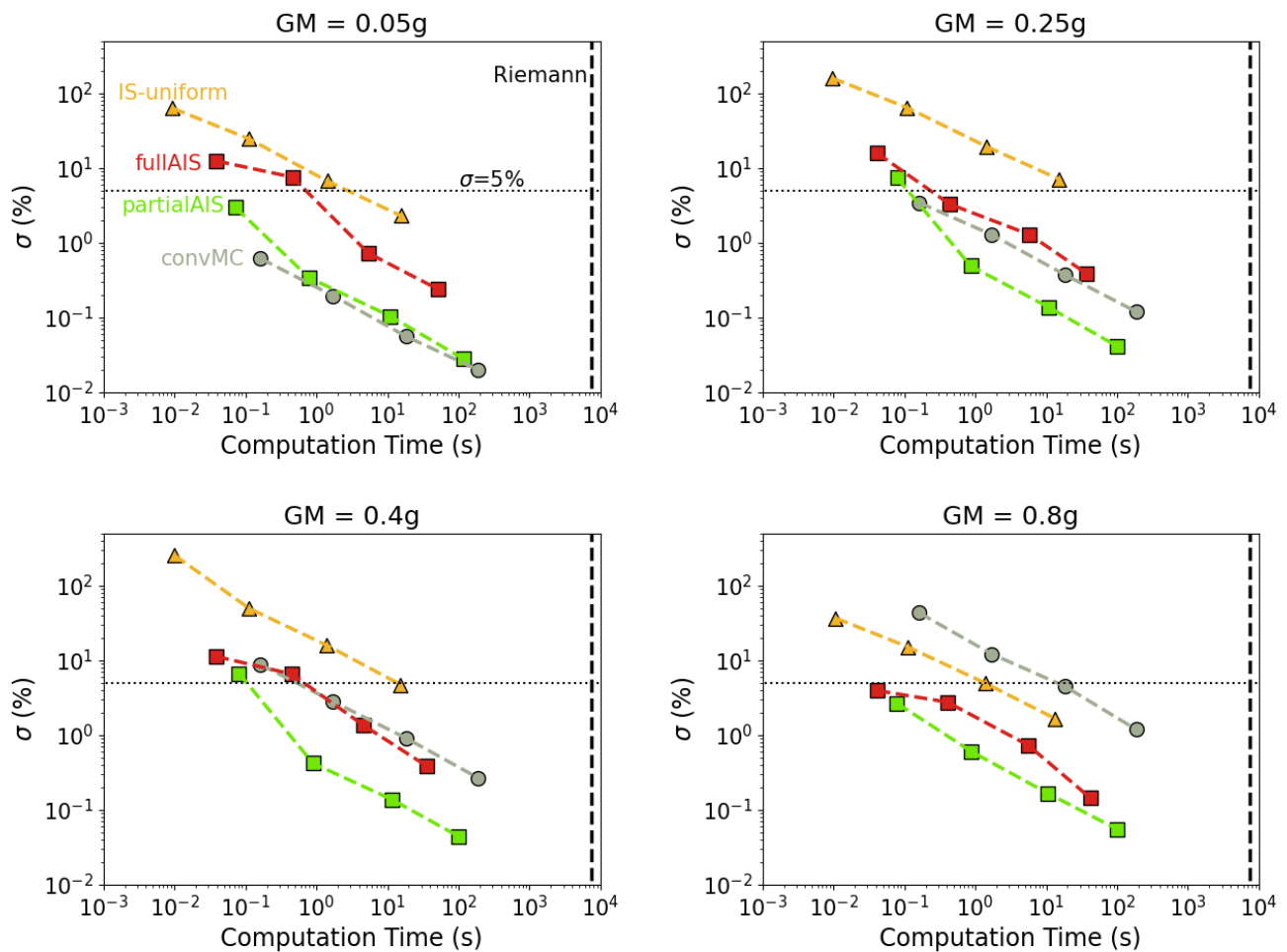
456 , which is the integral version of Eq. (1) and Eq. (5). Solving Eq. (39) posits computational challenges than the single source  
 457 problems because  $f_S(s)$  have the large jumps at  $i = 1, 2, \dots, N_s - 1$ . In addition, we introduce additional dependencies in the  
 458 multi-source case because  $m$  and  $r$  depend on  $s$ , further diminishing the VEGAS algorithm's effectiveness. Thus, we test  
 459 two AIS PSHA approaches for this case: 1) full AIS approach utilizing the Eq. (39) and 2) partial AIS approach, which is the  
 460 simple summation of single-source AIS PSHA curves.

461 We compared the computational performance of MC, IS, and full and partial AIS (Figure 9). Like the previous examples,  
 462 conventional MC is faster for low ground motions, e.g., 0.05 g, but not high ones. For example, to achieve the  $\sigma$  of 2 % at a  
 463 ground motion of 0.8 g, conventional MC, IS, and full and partial AIS take 73.8, 10.0, 0.77, and 0.13 seconds, respectively,  
 464 i.e., partial AIS is the most efficient algorithm,  $\sim 583, 79$ , and 6 times faster than conventional MC, uniform IS, and full AIS,  
 465 respectively.



**Figure 8.** Fault Source Example: The convergence of  $m$ ,  $r$ , and  $\varepsilon$  iterated IS densities density (red) derived in AIS algorithm ( $N = 1,000,000$ ) to marginal distributions of hazard deaggregation (black) at ground motion intensities of 0.1 g, 0.5 g, and 0.8 g.

466 As stated earlier, full AIS is less efficient than partial AIS because of the additional dependencies and jumps introduced  
 467 by  $f_S(s)$ . Figures S5 and S6 present the accuracy of both AIS approaches with different sample sizes. The contrast between  
 468 Figures S5 (c) S6 (c) shows that partial AIS has smaller error than full AIS even with fewer samples ( $150,000 < 500,000$ ).  
 469 We don't show the comparison for the partial AIS on ground motions greater than 0.8 g because fault B reaches a numerical  
 470 instability due to its low exceedance probability ( $< 10^{-12}$  /yr) (Figure S5).



**Figure 9.** Combined Sources Example: Standard deviation of the conventional MC (red), IS MC (blue), full AIS MC (black), and partial AIS MC (gray) estimates as functions of computation time at ground motion of 0.05 g, 0.25 g, 0.4 g, and 0.8 g

## 471 CONCLUSION

472 We proposed a novel computational framework for PSHA based on an implementation of the VEGAS algorithm. Through  
 473 comprehensive testbeds, we investigated the computational performance of this new approach, covering widely adopted  
 474 source types in PSHA practice: 1) areal source, 2) fault source, and 3) combined sources, including both areal and fault  
 475 sources.

476 We compared the proposed method to three existing computational frameworks: a) Riemann sum, which has exponentially  
 477 increasing computational costs for finer grid sizes; b) conventional MC, which requires a substantially long catalog with  
 478 synthetic earthquakes, particularly for large seismic intensities; and c) importance sampling (IS) with simple (uniform) IS  
 479 distributions.

480 Our findings indicate that AIS PSHA outperforms all other computational frameworks. AIS PSHA can dramatically reduce  
 481 computational times by factors up to  $> 10^5$  compared to traditional Riemann summation. AIS PSHA was also  $10^3$  faster than

482 the conventional MC while maintaining a 2% standard deviation of error. Additionally, AIS PSHA was up to 70 times faster  
483 than IS PSHA, demonstrating that the VEGAS algorithm can approximate optimal IS distributions quickly and well.

484 We showed that AIS PSHA requires a similar computation time for any ground motion, making its application to larger  
485 ground motions with low probability substantially more efficient than conventional MC. In PSHA practice, the computa-  
486 tional demands are dominated by large ground motions that need substantially more samples than lower ground motions.  
487 In contrast, AIS PSHA only requires a similarly low number of samples for all ground motion levels because it finds optimized  
488 IS distributions for each.

489 We also showed that AIS PSHA finds approximated deaggregation curves at no extra computational cost based on theoret-  
490 ical insights showing that optimal IS densities are equivalent to deaggregation distributions. We show empirically that the  
491 hazard deaggregation and iterated IS densities from AIS PSHA are fairly similar by comparing the statistical properties of  
492 the two distributions, e.g. K-S D statistics  $< 0.113$  and mean values differences of  $< 4.3\%$ .

493 Our study indicated that our AIS PSHA implementation works extremely well for point sources, where the magnitude  
494 and distance are independent random variables. We also applied the algorithm to problem employing finite-rupture model,  
495 and it still outperformed the pre-existing algorithms up to by factors of  $> 2,000$ . It appears to be less effective than point  
496 source example because it introduces a dependency between the magnitude and distance. Our implementation of the VEGAS  
497 algorithm uses independent random variables for the IS distribution, thus, making it less effective for the finite-rupture case.  
498 However, future implementations can also consider other versions of AIS algorithms where the variables are correlated;  
499 though computational demands for the AIS iteration would increase.

500 In the case of combined seismic sources, we proposed two strategies: 1) incorporating the source random variable into the  
501 AIS PSHA framework and 2) simple summation of AIS PSHA curves for individual sources. Both strategies outperformed  
502 the traditional methods up to by a factor of  $\sim 580$ . However, the second strategy was more efficient than the first one by a  
503 factor of 6. The first strategy added dependencies to the seismic hazard distribution, making the VEGAS less effective.

504 In sum, AIS can be applied to any PSHA computation, leading us to expect widespread application of the method.  
505 Specifically, we consider our proposed AIS PSHA to be significantly beneficial for large-scale projects that involve numerous  
506 logic tree branches and have extreme computational demands.

## 507 DATA AND RESOURCES

508 The source code for computing PSHA using the framework explained in this paper is available at  
509 [https://github.com/sehoung/ais\\_psha](https://github.com/sehoung/ais_psha). Figures were created using Matplotlib (Hunter, 2007) and Microsoft PowerPoint  
510 (<http://office.microsoft.com>; last accessed June 2024).

## 511 **DECLARATION OF COMPETING INTERESTS**

512 The authors have filed a patent application related to the method discussed in this paper. This application is relevant to the  
513 research presented and could represent a potential financial interest.

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## 633 APPENDIX A: VARIANCE OF MONTE-CARLO PSHA ESTIMATE

634 The second moment of  $\hat{\lambda}$ ,  $E[\hat{\lambda}^2]$ , can be expressed as:

$$\begin{aligned}
 E[\hat{\lambda}^2] &= E \left[ \left( \frac{\nu}{N} \sum_i^N I(X_i > a) \right)^2 \right] \\
 &= \frac{\nu^2}{N^2} E \left[ \left( \sum_i^N I(X_i > a) \right)^2 \right] \\
 &= \frac{\nu^2}{N^2} E \left[ \left( \sum_i^N I_i \right)^2 \right]
 \end{aligned}$$

635 Note that  $I(X_i > a)$  is denoted as  $I_i$  for simplicity in the following derivation. The above equation expands as follows:

$$\begin{aligned}
E[\hat{\lambda}^2] &= \frac{\nu^2}{N^2} E[(I_1 + I_2 + \dots + I_N)^2] \\
&= \frac{\nu^2}{N^2} E[I_1^2 + I_2^2 + \dots + I_N^2 + I_1 I_2 + I_1 I_3 + I_1 I_4 + \dots + I_1 I_N \\
&\quad + I_2 I_1 + I_2 I_3 + I_2 I_4 + \dots + I_2 I_N \\
&\quad + I_3 I_1 + I_3 I_2 + I_3 I_4 + \dots + I_3 I_N \\
&\quad \vdots \\
&\quad + I_N I_1 + I_N I_2 + I_N I_3 + \dots + I_N I_{N-1}] \\
&= \frac{\nu^2}{N^2} E \left[ \sum_i^N I_i^2 + \sum_{i \neq j}^{N^2-N} I_i I_j \right]
\end{aligned}$$

636 Here,  $I_i^2$  is the same as  $I_i$  since  $I_i^2$  also takes 1 if and only if  $X_i$  is greater than  $a$ . Also, the expectation operator can go inside  
637 the sum terms as it is a linear operator

$$E[\hat{\lambda}^2] = \frac{\nu}{N^2} \sum_i^N \nu E[I_i] + \frac{1}{N^2} \sum_{i \neq j}^{N^2-N} \nu^2 E[I_i I_j]$$

638 Here,  $\nu E[I_i]$  is equal to  $\lambda$ . Also,  $I_i I_j$  takes 1 if and only if both  $X_i$  and  $X_j$  exceeds  $a$ . Given that the ground motion samples  
639 ( $X_i$  and  $X_j$ ) are extracted independently,  $\nu^2 E[I_i I_j]$  is identical to  $\nu E[I_i] \nu E[I_j]$ , which is  $\lambda^2$ :

$$\begin{aligned}
E[\hat{\lambda}^2] &= \frac{\nu}{N^2} \sum_i^N \lambda + \frac{1}{N^2} \sum_{i \neq j}^{N^2-N} \lambda^2 \\
&= \frac{\nu}{N^2} N \lambda + \frac{1}{N^2} (N^2 - N) \lambda^2 \\
&= \frac{\nu \lambda - \lambda^2}{N} + \lambda^2
\end{aligned}$$

640 Hence,

$$\begin{aligned}
\text{VAR}[\hat{\lambda}] &= E[\hat{\lambda}^2] - E[\hat{\lambda}]^2 \\
&= E[\hat{\lambda}^2] - \lambda^2 \\
&= \frac{\nu \lambda - \lambda^2}{N} + \lambda^2 - \lambda^2 \\
&= \frac{\nu \lambda - \lambda^2}{N}
\end{aligned}$$

641 Note that  $E[\hat{\lambda}]^2$  is the same as  $\lambda^2$  since  $\hat{\lambda}$  is unbiased estimate.

## 642 APPENDIX B: VARIANCE OF IMPORTANCE SAMPLING PSHA ESTIMATE

$$\begin{aligned}
E[\hat{\lambda}^2] &= E \left[ \left( \frac{\nu}{N} \sum_i^N \frac{I(X_i > a) f_{M,R,\varepsilon}(M_i, R_i, \varepsilon_i)}{q_{M,R,\varepsilon}(M_i, R_i, \varepsilon_i)} \right)^2 \right] \\
&= \frac{\nu^2}{N^2} E \left[ \left( \sum_i^N \frac{I_i f_i}{q_i} \right)^2 \right]
\end{aligned}$$

643 Note that  $I(X_i > a)$ ,  $f_{M,R,\varepsilon}(M_i, R_i, \varepsilon_i)$ , and  $q_{M,R,\varepsilon}(M_i, R_i, \varepsilon_i)$  are denoted as  $I_i$ ,  $f_i$ , and  $q_i$  for simplicity in the following  
644 derivation. As derived in appendix A, the above equation expands as follows:

$$\begin{aligned}
E[\hat{\lambda}^2] &= \frac{\nu^2}{N^2} E \left[ \sum_i^N \left( \frac{I_i f_i}{q_i} \right)^2 + \sum_{i \neq j}^{N^2-N} \left( \frac{I_i f_i}{q_i} \right) \left( \frac{I_j f_j}{q_j} \right) \right] \\
&= \frac{\nu^2}{N^2} \sum_i^N E \left[ \left( \frac{I_i f_i}{q_i} \right)^2 \right] + \frac{\nu^2}{N^2} \sum_{i \neq j}^{N^2-N} E \left[ \left( \frac{I_i f_i}{q_i} \right) \left( \frac{I_j f_j}{q_j} \right) \right] \\
&= \frac{\nu^2}{N} \frac{1}{N} \sum_i^N E \left[ \left( \frac{I_i f_i}{q_i} \right)^2 \right] + \frac{1}{N^2} \sum_{i \neq j}^{N^2-N} \nu E \left[ \left( \frac{I_i f_i}{q_i} \right) \right] \nu E \left[ \left( \frac{I_j f_j}{q_j} \right) \right] \\
&= \frac{\nu^2}{N} E \left[ E \left[ \left( \frac{I_i f_i}{q_i} \right)^2 \right] \right] + \frac{1}{N^2} (N^2 - N) \lambda^2 \\
&= \frac{\nu^2}{N} E \left[ \left( \frac{I_i f_i}{q_i} \right)^2 \right] - \frac{1}{N} \lambda^2 + \lambda^2
\end{aligned}$$

645 Hence,

$$\begin{aligned}
\text{VAR}[\hat{\lambda}] &= E[\hat{\lambda}^2] - E[\hat{\lambda}]^2 \\
&= \frac{\nu^2}{N} E \left[ \left( \frac{I_i f_i}{q_i} \right)^2 \right] - \frac{1}{N} \lambda^2 + \lambda^2 - \lambda^2 \\
&= \frac{1}{N} \left( \nu^2 E \left[ \left( \frac{I_i f_i}{q_i} \right)^2 \right] - \lambda^2 \right)
\end{aligned}$$

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